

Bio-inspired wave breakers to reduce swell erosion in the Bay of Biscay using Computational Fluid Dynamics

Corentin Thomas, Victor Lieunard, Baptiste Oudon, Olivier Bain, Rejanne Le Bivic, Arnaud Coutu

► To cite this version:

Corentin Thomas, Victor Lieunard, Baptiste Oudon, Olivier Bain, Rejanne Le Bivic, et al.. Bioinspired wave breakers to reduce swell erosion in the Bay of Biscay using Computational Fluid Dynamics. Regional Studies in Marine Science, 2023, 62, pp.102965. 10.1016/j.rsma.2023.102965 . hal-04099641

HAL Id: hal-04099641 https://normandie-univ.hal.science/hal-04099641v1

Submitted on 11 Jan2024

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers. L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

1	
2	
3	
Л	
-	
5	
6	
7	
8	
9	Bio-inspired wave breakers to reduce swell erosion in the
10	Bay of Biscay using Computational Fluid Dynamics
11	
12	Corentin Thomas ^a , Victor Lieunard ^a , Baptiste Oudon ^a , Olivier Bain ^a , Rejanne Le
13	Bivic ^{a,} Arnaud Coutu ^{a b}
14	
15	^a Institut Polytechnique Unil aSalle Beguvais, 19 rue Pierre Waguet 60000 Beguvais, France
16	^b Corresponding author: <u>arnaud.coutu@unilasalle.fr</u> , 19 rue Pierre Waguet 60000 Beauvais, France
17	
18	
19	
20	
21	
22	LRH: Thomas, Lieunard, Oudon, Bain, Le Bivic and Coutu
23	RRH: Bio-inspired Wave breakers
24	
25	
26	
27	
28	
29	
30	
31	
32	
33 24	
34 25	
35	
30 27	
37	

38 Abstract

39

Coastal management is one of the main concerns of coastal countries. Sea level rise 40 caused by climate change can accentuate coastal erosion. One of its principal causes, 41 in France, is swell. Several solutions exist to limit the erosive effects of swell, including 42 wave breakers. Their main issue is that they protect a limited area and they can 43 increase the erosion downstream of the protected structures because of the cut in the 44 45 longitudinal transport. Therefore, existing wave breakers can't be a long-term solution to the problem of coastal setback. This paper is a proof of concept on the idea of 46 creating bio-inspired wave breakers that do not increase the erosion downstream but 47 it also shows that all the bio-inspirations are not appropriate. Two types of bio-inspired 48 wave breakers have been modelized on CatiaR2021X®. The first type is inspired by 49 the Torquigener albomaculosus' nest. The second type is inspired by the mangroves' 50 shape. Both shapes are known for slowing the seafloor current. The interaction 51 52 between the wave breakers and the swell in the Bay of Biscay has been simulated with 53 Xflow®2020. The experimental data needed to validate the simulations were provided 54 by the scientific literature. Because the results of the control simulations and of the simulation with existing wave-breaker shapes coincided, the results of the other 55 56 simulations have been considered as relevant too, even if test bench experiments will 57 be performed later to deepen these results. The simulations' results have shown that the nest-inspired shapes are increasing the erosion in the non-protected areas. On the 58 contrary, the mangrove-inspired wave breakers could be a possible improvement to 59 slow the coastal setback in the Bay of Biscay. This study aims to give first results about 60 the general interaction between geometrical shapes and swell, independently of the 61 materials of the wave breakers. 62

63 Keywords

64

Biomimicry, Erosion, Lattice-Boltzmann Method (LBM), Wall-Modeled Large Eddy
 Simulation (WMLES), Coastal research.

- 67
- 68
- 69
- 70
- 71 72
- 73
- 74
- 75
- 76
- 77
- 78
- 79

80 1. Introduction

81

With climate deregulation and global mean sea level (GMSL) rising, managing the 82 shorelines is one of the main concerns of coastal countries. GMSL has risen since 83 1900 by about 21 cm, including 7 cm since 1993 (Hay et al. 2015, Church et al., 2013). 84 Scientists' predictions cover 6 scenarios (Low to Extreme) that show that sea level 85 could rise from 30 cm to 250 cm by 2100 (Kopp et al., 2014; Lacroix et al., 2019; Miller 86 et al., 2013; Sweet et al., 2017). On low-altitude coasts, this phenomenon could 87 reinforce coastal erosion, and cause the loss of habitations, agricultural or natural 88 surfaces. In France, for example, 270 km of coastline are eroding at a rate superior to 89 50 cm/yr. (MilieuMarinFrance, 2020). One of the main causes of coastal erosion in 90 France, is swell (Aubie and Tastet, 2000; Sabatier et al., 2009; Migniot and Lorin, 1979). 91

This well-known problem has been studied for years. In France, between 1940 and 92 1950, three structures had already been built to decrease coastal erosion (Boulet et 93 al., 2018) and coastline management is a major concern of political institutions (French 94 95 Ministry of ecological transition, 2022). In the past century several solutions were tried, usually classified as soft or hard methods. Soft methods allow solutions to be 96 integrated into their environment. They are reversible and are cheaper in the short, 97 98 medium, and long term. According to Van Rijn (2011), they allow the sediment stock 99 to regenerate. Hard solutions include building wave-breaker or rockfill. These solutions are expensive but durable. One problem of wave breakers is that they only protect a 100 101 portion of the coastline and it induces hydraulic erosion downstream and upstream of 102 the protected areas (Boulet et al., 2018; Williams et al., 2018). This erosion increase 103 is a problem for large-scale coastal management because the aim is to protect the most eroded areas without eroding the other areas. This is why actual wave breakers 104 can't be a long-term solution to coastal setback problem until a solution will not be 105 found to make wave breakers that are not transferring erosion to other areas. Different 106 studies cover the optimum wave breaker parameters to reduce the erosion behind the 107 structure (Carpentier and Brosselard-Faidherbe, 1986; Delage, 1954; Massel and 108 Gourlay, 2000; Toyoshima et al., 1966) but few researches treat the question of finding 109 how to not increase erosion in non-protected areas. In France, where multiple 110 structures have been set up to reduce swell erosion, this recurrent problem doesn't 111 help to find large-scale solutions to the coastal setback. 112

France's coastline is 5,853 km long and faces 4 different seas and oceans with different 113 wave climates: North Sea, English Channel, Atlantic Ocean, and the Mediterranean 114 Sea. The English Channel and the Mediterranean Sea are exposed to wind sea and 115 short periods swell, while the Bay of Biscay, facing the Atlantic Ocean is exposed to 116 energic and moderate-to-long period swells (Dodet et al., 2019). This is explained by 117 the fact that swell energy depends on the wind speed and on the water surface in 118 contact with the wind, called the fetch. Studying a sand beach in the Bay of Biscay is 119 relevant to quantify the efficiency of the solutions to coastal erosion in an energetic 120 swell context. One of the possible solutions could be biomimicry. 121

Scientists used biomimicry and bio-inspiration for a long time to solve various problems. In 1903, the brothers' Wright plane was inspired by a pigeon flight (Cho and

Wood, 2016). Nowadays, shark skin is an inspiration for performant swimming suits 124 (Wen et al., 2014). It is important to understand that biomimicry isn't a goal but a way 125 to find innovative solutions to modern problems. In the wilderness, there are several 126 examples of natural or biological structures that reduce sea currents' impacts. The 127 discovery in 2012 of the Torguigener albomaculosus nest near the Japanese coast 128 has inspired multiple studies. This circular 2 m large structure reduces the speed of 129 the current in its centre (Kawase et al., 2013; Kawase et al., 2017; Matsuura., 2015). 130 Another inspiration could be the mangroves root system that has been known for a 131 long time for slowing sea current and consequently, helping in coastal stabilization 132 (Horstman et al., 2014; Hong Phuoc and Massel, 2006; Thom, 1967; Van Santen et 133 al., 2007). Both shapes could be an inspiration to create wave breakers. Because 134 seafloor current speed is directly linked to erosion (Everett et al., 2019; Migniot and 135 Lorin, 1979), the use of these shapes to create new types of immerged wave breakers 136 could reduce the current speed behind the structure without reflecting the energy to 137 the sides of the structure. In this case, wave breakers should not increase erosion of 138 non-protected areas. 139

This work aims to find new bio-inspired wave breaker shapes that could optimize the 140 erosion reduction behind the structure without eroding nonprotected areas in order to 141 be a solution to control coastal erosion. It is meant to be a proof of concept and 142 therefore start to quantify the interaction of these shapes and the swell. The location 143 has been taken as an example but the simulation data are not focused on one precise 144 location because the aim was not to study in detail one particular case. Wave-breakers 145 were 3D modeled in a submerged environment representing the Bay of Biscay area 146 and their impact on waves was studied by numerical simulation. Two bio-inspired forms 147 of wave-breakers were modeled based on the nest of Torquigener albomaculosus and 148 the root system of mangroves. These models were compared by numerical simulation 149 with usual wave-breakers. First, the modelling of the wavebreakers will be detailed in 150 this paper, then the results of the different numerical simulations, carried out with a 151 CFD tool as well as a Python program, will be explained in order to discuss the interest 152 of biomimicry for the protection against coastal erosion.2. Material and Method 153

154

155 2.1. Model development

In this study, a three-dimensional Computational Fluid Dynamics (CFD) software has 156 been used to calculate the effects of bio-wave breakers on sea currents and on swell. 157 It is based on the Lattice-Boltzmann Method (LBM). Simulations were made with a 158 free-surface model. It means that the fluid is in the lower part of the simulation and is 159 the only element considered (Fig. 1). The software resolves the Navier-Stokes and the 160 Boltzmann equations (Dassault Systèmes, 2021). The Navier-Stokes equations are 161 expressing the conservations of mass (Eq.1), of the momentum (Eq.2) and of the 162 energy (Eq.3). These are partial differential equations (Jamshed, 2015). 163

164
$$\delta \rho / \delta t + \nabla . (\rho \vec{V}) = 0 \quad : \quad Eq.1$$

165

$$\delta(\rho \vec{V}) / \delta t + \nabla (\rho \vec{V} \vec{V}) = -\nabla \rho + \nabla \Sigma + \rho \vec{g} \quad : Eq.2$$

168

$$\delta(\rho E)/\delta t + \nabla .(\rho E \vec{V}) = -\nabla .(P.\vec{V}) + \rho \vec{g}.\vec{V} + \nabla .\vec{q} + \nabla .\vec{q}_R : Eq.3$$

169

170 With ρ the density, *t* the time, *V* the eulerian speed, ∇ the divergence, Σ the stress 171 tensor, *g* the external mass-forces, *E* the total energy per mass unit, *q* the heat flow of 172 the thermal conductivity, and *q*_R the heat flow of the radiation.

173
$$f_{\alpha}(x + c_{\alpha}\delta x, t + \delta t) = f_{\alpha}(x, t) - 1/\tau * [f_{\alpha}(x, t) - f_{\alpha} e^{q}(x, t)] : Eq.4$$

174 With f_{α} the particle distribution function in direction, *x* the lattice node C_{α} the particle 175 discrete vectors, *t* the discrete times, δt the constant time step, τ the relaxation time 176 due to the fluid viscosity and f^{eq} the collision operator obtain with an order 2 Taylor 177 development of the Maxwell-Boltzmann distribution function.

The Boltzmann equation (Eq.4) is a distribution function of each group of particles. This 178 model assumes that the diffusion, or the motion of particles, is equal to the collision 179 (Heubes et al., 2013). LBM is a meshless particle-based method. The approach is 180 completely Lagrangian. It is one of the most powerful methods to calculate the 181 interaction between complex geometries and turbulent fluids (Satjaritanun et al., 2018). 182 It doesn't need to have a Sliding Mesh and a Multiple Reference Frame. The simulation 183 generates a Lattice element and organizes it in an Octree structure. It considers both 184 micro and mesoscopic processes in order to obtain correct macroscopic equations. 185 The LBM calculates a statistical distribution function with the variables, while 186 conserving mass, momentum, and energy (Eq. 1-3) (Chen and Doolen, 1998). 187

It uses a Wall-Modeled Large Edgy Simulation (WMLES) model. The hypothesis is 188 done that smaller turbulences are instantly and completely dissipating all the energy 189 they receive from the greater scale's structures (Grondeau et al., 2016). This increases 190 the efficiency by reducing the meshing operation and also the computational time 191 (Satjaritanun et al., 2018). The wall function used by the software to model the 192 boundaries is a unified non-equilibrium wall function. It means that the average velocity 193 is sensitized to the effects of the pressure gradient. It also uses the concept of two-194 layer-based to compute the sum of the kinetic energy of the turbulences (Dassault 195 Systèmes, 2021). The software chosen for the simulations was Xflow®2000, a 196 commercial LBM CFD solver. 197

198

In all the simulations, the values of some parameters have been kept. A 3D freesurface isothermal environment was used. The temperature variation was considered to be nil, therefore negligible even if it is not the case in the reality. The aim was to understand the interaction between the wave breakers shapes and a common swell. The gravity of the fluid was at 9.81 m.s⁻². This is due to the orientation of the spatial axes during the numerical simulation. The z-axis representing the altitude is usually oriented upwards, so gravity has a negative sign. The dimensions of the domain were

defined: 370 m long, 25 m high, and 350 m large (Fig. 1). All the water channel 206 boundaries were open except for the ground that was a free-slip rigid wall boundary. 207 The beach and the wave breakers were considered as free-slip rigid wall boundaries 208 too (Fig. 1). The mesh was 0.25 m large around the wave breakers and 1 m large 209 everywhere else. This size was conditioned by the software and by the computer. If 210 the mesh size was smaller, the software couldn't generate the Lattice domain. The fluid 211 was seawater with a temperature of 20°C and a density of 1024 kg.m⁻³ (Vandermeirsch 212 et al., 2012). 213

214 Fig.1 (Small column size)



215 L216 2.2. Parameters of the simulation

For this study, the action of wave breakers was simulated in the Bay of Biscay in 217 France. According to Migniot and Lorin (1979), the swell is more powerful in the Atlantic 218 than in the Mediterranean Sea, therefore selecting the Atlantic Ocean as a study area 219 is interesting for the quantification of the maximum effect of the swell. The goal was to 220 221 simulate a similar morphology as the beach on the coast, near Biscarrosse (France) (44° 21' 3.55" N 1° 9' 58.424" W). In order to achieve it, the bathymetric data from 222 January 2021 of the European Marine Observation and Data Network (EMODnet) has 223 been used to design the beach profile with CatiaR2021X® (Fig.2 A, B). The beach 224 profile was inspired by these data and extrapolated to get a 350 m large structure. This 225 could not be considered as a realistic beach but the goal of this study is limited to 226 understand the interaction between the wave breakers and the swell. The longshore 227 bar was considered as fixed in this study because the simulation duration was of 45 s. 228 For each simulation and each wave breaker, the depth at the top of the wave breaker 229 was 1.5 m and the mean distance with the coast was the same. A study considering 230 231 shape complexity and mechanical characteristics of each wave breaker could be interesting but is out of scope for this work. 232

233

234 Fig.2 (Small column size)



As explain in Brivois et al., in 2012,"The Aquitanian Coast (SW France) is an 236 approximately 250 km long straight low coast bordered by high aeolian dunes". 237 Prederos et al (1996) describes the sediments of the study area as medium to fine 238 239 sands with a grain size ranging from 200 to 400 mm. According to Wright and Short 1984 classification, the majority of Aquitaine beaches are intermediate beaches with 240 Transverse Bar and Rip (TBR), or low tide bar/rip as defined by Masselink and Short 241 (1993). During spring tides, in the study area, the sea level rises from 0.3 m to 5 m 242 (Butel et al., 2002), classifying these beaches as mesotidal. 243

As explained by Castelle and Bonneton in 2002, tidal currents are largely negligible compared to those induced by the swell. However, they were calculated by Sénéchal *et al.*, in 2009 and have a speed of 0.2 m.s-1 from West to East during the flood and from East to West during the ebb.

The wave regime has a mean annual significant wave height of 1.36 m and a mean period of about 8 s with a strong seasonal dependence linked to the NAO (Butel *et al.* 2002; Dupuis *et al.*, 2006; Sénéchal et al., 2009; Carles *et al.*, 2011; Le Cozannet *et al.*, 2011et al. 2009, Brivois *et al.*, 2012). Brivois *et al.*, (2012) describe the wave conditions as evolving from low energy waves during the summer oriented NW to very high energy waves during winter storms generating west to NW swells.

Regarding the wave breakers characteristics, several studies cover the optimum parameters (Carpentier and Brosselard-Faidherbe, 1986; Delage, 1954; Massel and Gourlay, 2000; Toyoshima *et al.*, 1966). For example, the efficiency of a structure is linked to its depth and its distance to the coast (Delage, 1954). These parameters are *D* the distance between the wave breaker and the coast, *L* the length of the structure, *H* its height, *p* the depth where the wave breaker is located, λ the swell wavelength, and *T* the wave period.

262

263	Table	1

Parameter	Value	References
D	225	Massel and Gourlay, 2000
	m	
L	285	Toyoshima <i>et al</i> ., 1966
	m	
Н	3.55	Delage, 1954
	m	
р	4 m	Found in the simulation
λ	102	Carpentier and Brosselard-Faidherbe, 1986
	m	
Т	8.0 s	Carpentier and Brosselard-Faidherbe, 1986

264

265 2.3. Existing wave breakers

The first wave breakers modelized were inspired by existing shapes like trapezoidal or 266 elongated tubes. Both shapes were found in the literature. Trapezoïdal shape for 267 immerged wave breaker is common in mathematical articles, bench tests, and real 268 beach analyses (Hornack, 2011; Repousis et al., 2014). Authors assume that other 269 types of existing wave breakers could be used, but the main goal of this study was only 270 to suggest new bio-inspired shapes for wave breakers and not to make an exhaustive 271 study on wave breakers. This geometrical shape is easier to create and to settle on 272 the seafloor than more complex shapes. Elongated tubes are another wave breaking 273 shape which is induced by the process of creation of the wave-breaker. They are 274 Multipurpose Artificial Reefs (MPAR) and are used to generates a surfable wave break 275 type as well as reducing the swell energy. A geotextile tube is set up, then filled with 276 sand before being settled down on the seafloor. It is easier and cheaper to install than 277 concrete, and its shape is more continuous and regular than a rockfill. They could be 278 damaged by a rupture of the geotextile. This shape has been studied by different 279 studies with Computational Fluid Dynamics and beach studies (Hedge, 2010; Isebe et 280 al, 2008). Both shapes have been modelized on the 3DEXPERIENCE Platform with 281 CatiaR2021X®. 282

284 2.4. Bio-inspired wave breakers

285 The first inspiration for designing bio-inspired wave breakers was the *Torquigener* albomaculosus' nest. This fish was observed in the Pacific Ocean, near the Japanese 286 coast (Kawase et al., 2013). Different nests have been found between 10 and 30 287 meters deep. The 10 cm fish takes less than 10 days to create a circular 2 meters large 288 shape in the sand. It is composed of two concentric parts (Fig.3 C). The outer ring with 289 24 to 28 cm high peaks, 12 cm apart from each other. The inner circle is a small 290 depression with an irregular pattern (Kawase et al., 2017). All the nests have usually 291 the same outer ring. There is on the outside part a bigger dune and on the inside part 292 a smaller one. This nest is exposed to currents going between 0.8 and 1.0 m.s⁻¹. The 293 nest could cause a slowdown of the current because of its particular geometrical 294 pattern. This current could be broken by the greater ripple and then be uniformed in 295 the centre of the nest by the smaller one (Kawase et al., 2013; Kawase et al., 2017; 296 Matsuura., 2015). 297

The nest has been modelized on the 3DEXPERIENCE Platform with CatiaR2021X®. 298 2D sketches and multi-section solids were used to create a part of the nest. Then, it 299 was duplicated 24 times to get a 3D model of the entire nest (Fig.3 C). Following this 300 step, different types of wave-breaker were modelized. In order to do so, models of 301 existing shapes as a trapeze and three elongated tubes were done (Fig.3 A, B). 302 Following this logical path, shapes inspired by the nest were created. Each model 303 added another characteristic of the nest. The first one (Fig.3 D) was straight and kept 304 the order of the two ripples; the taller one before. The second one (Fig.3 E) was still 305 straight but put an angle between every two dunes. The third one (Fig.3 F) added the 306 circular pattern by being built with guarters of the nest. The fourth one (Fig.3 G) took 307 larger parts of the nest; it was made with thirds of the nest. The fifth one (Fig.3 H) 308 continued this logic with halves of the nest. The sixth one (Fig.3 I) took quarters of the 309 nest but the taller dunes were always facing the sea. From the third to the sixth, wave 310 breakers had a sinusoidal shape. The last one (Fig.3 J) was the merge of two shapes. 311 The base was a trapeze-shaped wave-breaker and the top was a straight bio-inspired 312 shape (Fig.3 D) All these shapes were rescaled to have the same height and the same 313 length given in table 1. 314

315

316 Fig.3 (Small column size)



The second inspiration was the mangroves. Horstman et al. (2014) described 318 mangroves as an ecosystem located on subtropical tidal coasts. They are composed 319 of trees and shrubs from distinct species which are exposed to waves and flooding. 320 They form, with a coral reef generally, a buffer between land and sea, contributing to 321 wave attenuation (Horstman et al., 2014; Hong Phuoc and Massel, 2006) and 322 consequently, helping in coastal stabilization(Anthony, 2004; Van Santen et al., 2007). 323 Thereby, the question is, can mangrove-inspired wave breakers be effective in 324 325 reducing coastal erosion? Mangroves are constituted of varied species of plants that interact differently with the swell and can be more or less effective to protect the 326 coastline (Hong Phuoc and Massel, 2006). One of the species with the highest wave 327 height reduction is the Rhizophora (Fig.4 A) (Phan et al., 2014; Lee et al., 2021). 328

329 In order to test the capacity of mangroves to act as wave breakers, 3D models have been created on the 3DEXPERIENCE Platform with CatiaR2021X. Two main types of 330 models have been tested, following a symmetric geometry or a chaotic geometry. The 331 symmetric geometry had three axes separated by an angle of 60° with a radius of 7.78 332 333 meters and a height of 3.55 meters. Different roots' densities have been tested with 3, 5, 8 and 10 vertical roots per axis (respectively Fig.4 E,D,C,B). The chaotic structure 334 followed a semi-sphere geometry with a radius of 3.71 meters and a height of 3.55 335 meters (Fig.4 F). The tools used most of the time on CatiaR2021X to create these 336 models were the extrusion, used to create the principal structure, and the ridge, used 337 to create the branch attached to the main structure. Each model has been duplicated 338 to build rows to form wave breakers. There are a 15m distance between each centre 339 of mangrove so that there are 0.5 m of gap between two shapes. The aim of testing 340 these shapes is not to study the technical properties of the wave breakers but to reveal 341 if their interaction with the swell can reduce the energy of the seafloor currents. For the 342 same reason, the material used to create these shapes is not considered. 343

344

345 Fig.4 (Small column size)



346

347 2.5 Agents defining erosion

348

Hjulström (1935) defined grain size and current velocity as important factors in sedimentary movement. Subsequently, in the 1970s, sediment transport modelling was initiated, resulting in the development of several formulae that describe longshore drift - a critical aspect of sedimentary stock movements - and lead to erosion. These formulae were proposed by Longuet-Higgins (1970), Komar and Inman (1970), Komar (1979), and Cerc (1984).

355

Whether it is a reflective or diffractive shoreline, as defined by Masselink and Short (1993), sedimentary transport important factors are: i) morphology (beach barred or unbarred), ii) slope (beach and nearshore slope), iii) grain size, iv) wave height and angle and v) currents. With the tidal current's exception, the currents are derived from the height and angle of the swell (drift current, swash and backwash current). To limit erosion, it is necessary to reduce: the wave height and/or the wave angle and/or the currents.

363

The model 's input parameters are: i) granulometry (between 0.10 and 0.20 mm) and 364 ii) bottom currents defined by Migniot and Lorin study in 1979 (1.4 m/s maximum). The 365 grain size is not simulated. It is used in this study to interpret simulation results. For a 366 particle size between 0.10 and 0.20 mm, a current of 1.0 to 5.0 cm/s speed will 367 368 transport the particles (suspension, creep and saltation) with a 20 cm/s speed. The simulations objectives are: i) to observe and understand the interaction between 369 currents/swells and new forms of bio-inspired wave breakers, and ii) to observe, if 370 possible, the variations of current speed with the implementation of wave breakers. 371 The drift current will not be considered in these simulations, nor the wave height due 372 to computational cells size. 373

374

2.6. Location of the observation points

376 A simulation with all the characteristics identified above has been created. 56 observation points (or probes) were located as eight lines of seven on the seafloor. 377 They are fixed points that identify the nearest cell to their locations and write out the 378 cell values with a certain frequency. For example, in the simulations, each point 379 recorded the value of the speed every 0.00345 seconds (290 Hz). The four first lines 380 (points 1 to 28) are placed on a gentle slope from the entrance of the simulation to the 381 edge of the longshore bar, a submarine dune (Fig.2). The fifth line (point 29 to 35) was 382 just behind this bar. The sixth one (points 36 to 42) is in the surf zone. Finally, the two 383 last lines (points 42 to 56) are on the floor of the swash zone (Fig.2). 384

Then, two situations were studied: at ebb tide and flood tide. These were the control simulations. Each observation point gave more than 13,000 speed values for each simulation. Then the simulation was kept the same as were implemented the wave breakers, one at a time. For every shape, two simulations for both tide were done. In addition, the Python program was used to compare these simulations with their controls. With this method, it was possible to characterize the efficiency of the wavebreaker to slow down the current and consequently the swell erosion.

- 392
- 393 3. Results
- 394
- 395 3.1. Control simulations

The first thing to do was the control simulations. There were two with the average swell 396 at both tides. When the depth is lower than one and a half time the height of the wave, 397 the wave collapses. This always happened just before the longshore bar (Fig.2). After 398 this, the sea seemed smoother but the current was still observed with a maximum 399 average speed of 1.38 m.s⁻¹ on the seafloor. This would be enough to move sand 400 particles easily. The fastest currents seemed to stay at the top of the water. This was 401 a good thing because waves were already broken by the longshore bar and there was 402 still particle displacement, which implied positioning the wave breakers upstream of 403 the bar so that the waves collapsed earlier. 404

Regarding the seafloor current speeds, at flood tide, the speeds were between 0.55 and 0.45 m.s⁻¹ at the three first points lines. The current accelerated around the longshore bar with 0.7 m.s⁻¹ on the fourth line and 0.85 m.s⁻¹ on the fifth line. Then, the seafloor current slowed with a 0.80 m.s⁻¹ speed at the sixth line and a 0.62 m.s⁻¹ speed at the seventh line. The sea level reached sometimes the eighth line but it registered an average speed of 0.05 m.s⁻¹. The average speed of all points was 0.52 m.s⁻¹.

At ebb tide, the speed decreased on the first three lines from 0.46 m.s⁻¹ to 0.26 m.s⁻¹. Then, the current accelerated to reach 0.50 m.s⁻¹ on the fourth line and 0.72 m.s⁻¹ right behind the longshore bar, on the fifth line. At the sixth line, the maximum seafloor current speed was reached with 1.38 m.s⁻¹. Then, the speed decreased to 0.76 m.s⁻¹ on the seventh line. The sea level didn't reach the last line so that the last points didn't register any value. The average value for all the points was 0.63 m.s⁻¹.

418 3.2. Existing wave breakers

The second step was the simulation of both existing models with the average swell. 419 The depth of the top of the structure was 1.5m. The wave breakers had a similar 420 behaviour (Fig.5). At flood tide, there was on the third line of points a slowdown 421 between -0.4 and -0.35 m.s⁻² that kept the speed near 0.0 m.s⁻¹. The negative 422 acceleration means a mean reduction in the acceleration of the current measured at 423 the observation points. The fourth observation point line showed the same changes 424 but with smaller intensity. From probes line n°5 to n°7 (probes 29 to 49 represented in 425 Fig.5), there were no major changes between the control and the simulations. The last 426 line showed that the sea level had not reached all the observation points. At ebb tide, 427 the behaviour of the current was completely different. The biggest slowdown in both 428 cases was -0.2 m.s⁻² and some accelerations were over +0.3 m.s⁻² for the elongated 429 tube and +0.7 m.s⁻² for the trapeze. These maximum accelerations were located on 430 the fifth line, right after the longshore bar. The average change was -0.091 m.s⁻² at 431 flood tide and +0.034 m.s⁻² at ebb tide with the trapeze. With the elongated tubes, it is 432 -0.038 m.s⁻² at flood tide and +0.001 m.s⁻² at ebb tide. 433

434

435 Fig.5. (Full page width)



436

437 3.3. Nest-inspired wave breakers

The third step was the simulation of the nest-inspired wave breakers with the average swell. Both straight wave-breaker created little accelerations (e.g. +0.1 m.s⁻²) and a general slowdown at flood tide. These results changed drastically at ebb tide. There were some accelerations of +0.25 m.s⁻² around the third line (just behind the structure) and around +0.15 m.s⁻² next to the fourth and fifth line, around the longshore bar. The four sinusoidal wave-breaker had different results. They all presented some zones with

accelerations (e.g. +0.4 m.s⁻² for the shape with thirds of nest at flood tide) and other 444 areas with slowdowns (e.g. -0.19 m.s⁻² on the same model at ebb tide). The location 445 of these zones and the values were depending on the shape and the tide. None of 446 them showed a general slowdown at both tides. The last simulation with the nest shape 447 was the one with the hybrid wave-breaker (J on Fig.3). It was the only shape that 448 generated a general acceleration at the fifth and sixth observation point lines, right 449 behind the longshore bar. The average speed value was 0.56 m.s⁻¹ with the wave-450 breaker and 0.53 m.s⁻¹ without it. At ebb tide, there were still consequent acceleration 451 zones (e.g. +0.5 m.s⁻²) but no more important slowdown areas (Fig.6). The average 452 speed (0.66 m.s⁻¹) was superior to the control average value (0.63 m.s⁻¹). To sum up, 453 no model, with these parameters, generated a global slowdown of the current on the 454 seafloor. 455

456



457 Fig.6. (Full page width)

458

459 3.4. Mangrove-inspired wave breakers

The fourth step was the simulation of the mangrove-inspired wave breakers with an 460 average swell. On the first hand, simulations with two rows of a symmetrical model 461 have been done (Fig.4 B). At flood tide, only four observation points showed an 462 acceleration over +0.1 m.s⁻². They were all on the sides of the second and third points 463 lines, just next to the structure. Seven points showed a slowing of -0.5 m.s⁻². The 464 average seafloor speed changed from 0.53 m.s⁻¹ in the control to 0.44 m.s⁻¹ in this 465 case. At ebb tide, the maximal acceleration was around +0.2 m.s⁻² at point 32, just 466 behind the longshore bar. 42 points showed a slowdown and the average speed 467 changed from 0.63 m.s-1 in the control to 0.55 m.s⁻¹. 468

On the second hand, simulations with two rows of a chaotical model have been done.At flood tide, the first four observation points lines showed an acceleration. Five points

had a change over +0.1 m.s⁻² and two were over +0.2 m.s⁻². The four last lines put in evidence a slowdown that reached -0.54 m.s⁻² at the seventh line. The average seafloor speed changed from 0.53 m.s⁻¹ in the control to 0.42 m.s⁻¹ in this case. At ebb tide, only the seven points of the fifth line were showing an acceleration. This maximum value was at +0.08 m.s⁻² and the biggest slowdown was at -0.73 m.s⁻² at point 37, on the sixth line. The average speed changed from 0.63 m.s⁻¹ in the control to 0.44 m.s⁻¹.

477

Then, several roots' densities have been tried to determine if there was a link between 478 density and efficiency. At flood tide, the simulation with two vertical roots per axis 479 showed accelerations at the centre of each observation point line. The average speed 480 was higher than the control simulation with a value of 0.58 m.s⁻¹. The other simulations, 481 with four, eight and ten roots per axis, showed more or less the same behaviour. The 482 first four lines showed an acceleration. Two points had a change over +0.1 m.s⁻² and 483 were at +0.2 m.s⁻². The three following lines put in evidence a slowdown that reached 484 -0.50 m.s⁻² at the seventh line. The eighth line showed an acceleration between +0.1 485 and +0.0 m.s⁻². The average speed was at 0.53 m.s⁻¹ in the control and 0.44 m.s⁻¹ in 486 all the simulations. 487

At ebb tide, on the one hand, the behaviour was the same with 2, 8 or 10 roots per 488 axis. The average speed was around 0.55 m.s⁻¹ with a control at 0.63 m.s⁻¹. There 489 were less than ten observation points that showed a non-zero acceleration with only 490 five observation points over +0.1 m.s⁻². Only the fifth observation line, just behind the 491 longshore bar put in evidence a general acceleration. The other points registered a 492 slowdown between -0.05 and -0.31 m.s⁻². On the other hand, with 4 roots on each axis, 493 the behaviour was completely different. There were only two points with accelerations 494 of 0.11 and 0.08 m.s⁻². Moreover, the sixth line showed a slowdown around -1.3 m.s⁻² 495 (Fig.10 B). The average speed was around 0.22 m.s⁻¹. One important thing was that 496 the maximum speed was at 0.38 m.s⁻¹ when it was 1.4 m.s⁻¹ in the control (Fig.7 B). 497

- 498
- 499 Fig.7. (Full page width)



For the fifth step, different numbers of consecutive lines of the chaotical model have 501 been tested to determine if there was a link between this parameter and the efficiency 502 503 of the wave breaker. Each line's centre was about 15 m apart, which resulted in a gap 504 between lines of 0.5 m. This is a gap similar to the one between each mangrove in the line. At flood tide, the general behaviour was the same but the values were changing 505 depending on the number of lines. There were a few accelerations over + 0.1 m.s⁻¹ and 506 one over +0.2 m.s⁻² at point 14. On the fifth line of points, the only simulation to show 507 an acceleration was the one with three rows of mangroves. The simulation with two 508 rows registered the lowest or second-lowest seafloor current speed at every point 509 510 except at points 14, 23, 53 and 55. The average speed was at 0.44 m.s⁻¹ with one row, at 0.43 m.s⁻¹ with two rows and 0.49 m.s⁻¹ with three rows (Fig.8 A). 511

At ebb tide, the general behaviour was still the same between the three simulations 512 but the values were changing depending on the number of rows. Only a few points 513 registered an acceleration but on the fifth line of observation points, there was an 514 acceleration that could go up to $+ 0.2 \text{ m.s}^{-1}$ when there were three mangroves' rows. 515 In general, the average seafloor current speed with two rows of mangroves was the 516 lowest. The only exception was between points 42 and 48, where the lowest value was 517 in the one-row case. The average speed was at 0.63 m.s⁻¹ in the control, at 0.51 m.s⁻¹ 518 with one row, at 0.44 m.s⁻¹ with two rows and at 0.58 m.s⁻¹ with three rows (Fig.8 B). 519

520

521 Fig.8. (Full page width)



523 Finally, different current orientations were tested on the most efficient mangrove-524 inspired wave breakers. The goal is to observe if different current orientations will 525 change the impact of the wave breakers on erosion. For this purpose four different 526 orientations were tested for the flood tide. An angle of 30, 45, 60 and 75 degrees to 527 the normal flow was chosen. A control simulation was run for each of these angles to 528 compare the results. The average velocity of the control simulation and the breaker 529 simulation is shown in Table 2

530

531 Table 2

Flow angle	75°		60°		45°		30°		0	
Simulation	Control	Wave- Breakers								
Mean Velocity (m.s-1)	0,83	0,74	1,05	0,94	0,58	0,56	0,57	0,54	0,52	0,43

532

The simulations performed with an angle will result in an increase of the water velocity 533 at the ground level. The behaviour of the water deceleration is different for the 534 simulation with orthogonal current than for the simulations with an angle of current. But 535 this behaviour is similar between the different simulations with a current angle. The 536 majority of the deceleration is located on the third line (Fig.9), ranging from a 537 deceleration of 1.8 m.s⁻¹ for the 60° simulation to 0.55 m.s⁻¹ for the 30° simulation. 538 From the 4th to the 5th line an increase of the speed oscillates between 0 and 0.4 m.s⁻ 539 ¹. For the current at 45° a decrease in velocity from 0.24 to 0.13 m.s⁻¹ for the 6th and 540 7th line appeared. Whereas the simulation at 75° leads to an increase in velocity from 541 0.35 to 0.25 m.s⁻¹ 542



545 3.5 Validation of numerical simulations

544

The experimental data needed to validate the simulations was found in the scientific literature. The results of the control simulations and simulation with existing wave breakers have been compared with experimental data. If it validated the first simulations, it would make the other simulations with new types of submerged wave breakers more credible.

Firstly, the results of the control simulations have been compared to check the validity 552 of the interaction between the beach and the wave function. The main criteria to 553 compare natural waves and simulated waves were the breaking phase and location. 554 According to Larson and Kraus (1992), the waves breaking zone is located around the 555 longshore bar when there is one. Migniot and Lorin (1979) explain this with the fact 556 that the waves tend to collapse when the water depth is smaller than 1.5 times the 557 wave crest height. This behavior was present in the simulations. The waves collapsed 558 just before the longshore bar (Fig. 2). At this location, the slope made the depth lower 559 560 than 3 meters and the wave crests were 2 meters high. It was natural that waves broke at this location. 561

Secondly, the interaction between the fluid and existing wave breakers has been 562 compared with experimental data. The first point that was common in most papers is 563 that the wave-breaker slows the current right behind it (Balouin et al., 2016; Boulet et 564 al., 2018; Rasidah et al., 2010; Williams et al., 2018). At flood tide, this behavior was 565 observed with both existing wave breakers (Fig. 5A) (points 16-19 and points 23-26). 566 The slowdown was maximum just behind the center of the structure (points 16-19). 567 According to Balouin et al. (2016), the submerged wave-breaker protects the coastline 568 behind it but tends to parallelize the shoreline to the structure by eroding the changes 569

in the shoreline direction. Balouin et al. (2016) explain it by a change in the seafloor 570 currents directions that are deflected by the structure. Rasidah et al. (2010) have 571 noticed the settling of turbulences on the edges of the structure that could be due to 572 the diffraction of the waves' energy. In the simulations, the observation of the velocity 573 vectors showed that the current was deflected from the centre of the structure to its 574 edges and then to the coast. This caused turbulences on the edges that could have 575 been interpreted as side effects if it was not present in experimental data. In addition 576 to the direction of the velocity vectors, this effect was seen in the values of these 577 vectors at both tides (Fig. 5A,B) (Points 8, 14, 15 and 21). Because of the experimental 578 validation of the control simulations and the one with existing submerged wave 579 breakers, the results of the other simulations could be considered as first results on 580 the use of bio-inspired submerged wave breakers. 581

582 4. Discussion

583

584 4.1 Existing wave breakers

This study aimed to find a solution to one of the main problems of wave-breaker: the 585 increasing erosion of non-protected areas. When a structure is settled, it protects some 586 areas by decreasing waves energy and slowing the current on the seafloor but it 587 reflects some of the swell energy. This reflection increases the current at another place 588 and that will enhance the erosion in most cases (Williams et al., 2018). This reflection 589 was present in the simulation. The speed vectors were changing direction when they 590 had a contact with the wave-breaker and there were accelerations on the sides of the 591 structure (EMODnet., 2020). The followed method allows verifying the risk of eroding 592 non-protected zones by comparing the speed of seafloor currents with and without a 593 wave-breaker at different places. If a structure caused a general slowdown in current 594 around and after it, it means that the energy was dissipated and not reflected which 595 would not increase the erosion in the non-protected areas. On the other hand, if a 596 model generated accelerations in some zones, it would increase the erosion in these 597 places and consequently cause a destabilization of the sediment dynamics. 598

599 The existing models were slowing the current at flood tide, and thus were decreasing the erosion. There were some little accelerations at some observation points but 600 nothing that would risk causing some major changes. At ebb tide, however, the 601 accelerations areas (+0.3 or +0.7 m.s⁻²) were destabilizing the sedimentary cell. These 602 accelerations would cause the increase of erosion in several areas and this could 603 604 change the morphology of the beach. For example, the huge acceleration just behind the longshore bar could move faster this structure in the medium or long term. Even if 605 this bar was considered as fixed because of the 45 seconds duration of the simulation, 606 607 it is important to understand that this structure is moving naturally with the tides and the swell. These results went in the direction of previous studies which claimed that 608 building offshore structures as wave-breaker increase the erosion of some areas. 609

610

611 4.2 Nest-inspired wave breakers

The nest-inspired wave breakers had diverse behaviour and could cause a slowdown 612 of the seafloor current in some areas, depending on the tide. However, every model 613 was causing accelerations in some areas and the average slowdown was not 614 consequent. These accelerations could go up to +0.5 m.s⁻². This value, compared to 615 the size of sand particles using the Hjulström diagram (Hjulström, 1935), was very high 616 and even with a base value of 0.0 m.s⁻¹, the acceleration would be enough to erode 617 and move sand particles. No model had shown a general slowdown at the ebb and 618 flood tide. Finally, these really large structures (285 m long) were not causing an 619 average slowdown that would be relevant to a use in real conditions. For these 620 reasons, none of the nest-inspired models tested seemed to be a solution to manage 621 the general coastal setback. 622

623

4.3 Mangrove-inspired wave breakers

The mangrove-inspired immerged wave breakers had diverse behaviours but they seemed way more interesting than the nest-inspired ones. They all showed a general slowing that went up to -0.40 m.s⁻² at ebb tide. Above all, these models presented no zones with important accelerations. At flood tides, there had been some small accelerations (less than +0.2 m.s⁻²) between lines one and five. A general slowdown of the seafloor current meant decreasing erosion and increasing sedimentation. Therefore, these results could show a solution to the swelling erosion.

To estimate how efficient mangrove-inspired wave breakers could be, two parameters 632 were were considered : the number of rows and the density of roots (Fig.7 and Fig.8). 633 Setting up two lines of artificial mangroves seemed to be the best solution (Fig.8) 634 635 because it caused the most important slow down at both tides for most of the observation points. Two consecutive rows would cause the most important decrease 636 to the swell erosion and could slow coastal setback in the Bay of Biscay. Changing the 637 number of vertical roots had not caused an important change at flood tide but at ebb 638 tide, with the four-roots-per-axis model, the seafloor current speed was always under 639 0.4 m.s⁻¹ (Fig.10). Because this model didn't seem to create accelerations zones 640 around the structure and generate a global seafloor current slowdown, it could be the 641 most efficient to answer the problem of coastal setback. 642

By dissipating the energy of the wave, the impact on coastal sediments should be 643 decreased in a meaningful proportion, decreasing the general coastal erosion. The 644 efficiency of vegetation in dissipating wave energy, especially in mangrove settings, 645 was established by numerous studies (Dalrymple et al., 1984). More recently, Zhang 646 et al. (2020) applied Dalrymple's model to estimate how the wave attenuation was 647 changing depending on the mangrove's structure. Overall a decrease in wave height 648 is observed in any case of mangrove situation, which is corroborated by other studies 649 (Phan et al., 2014). Most of those studies (Mendez et al., 2004) occur in a study area 650 that is similar to the conditions of our study: fine sediments, depth below 10m, and 651 similar wave speed. However, it is important to wonder if some scale effects can also 652 be at play and impact sediment erosion around the wave breaker. It is important to be 653 aware of the turbulence that could be generated by the roots-like structure at a smallest 654 scale than the experimental scale used in the study. Recent papers on sediment 655

transport in vegetated channels show that, especially in mangrove-like areas, local 656 turbulence can be generated by the presence of roots study (Yang and Heidi, 2018). 657 These turbulences could create local vortex, able to lift up particles and affect the 658 bedload. This effect should be considered for former experimentations in this particular 659 setting. Other recent studies tend to induce that erosion in sand environments in 660 mangrove forests should consider near-bed turbulent energy instead of general 661 velocity (Norris et al., 2021). All those physical mechanisms should be considered 662 the impact of mangrove-like when studying wave breakers and before 663 implementation. 664

665 Fig.10. (Full page width)



666

667 4.4 Efficiency comparison

In order to compare the efficiency of the models, the average speeds of all simulations 668 have been placed in the Hjulström-Sundborg diagram. The average particle size in the 669 Bay of Biscay is 0.15 mm. This means that even if the size of the particles changes 670 along the coastline and along the profile of the beach, this size helps to understand the 671 general behaviour of the particles. All the points, no matter the tide or the model used, 672 were placed in the green circle (Fig.11 A). The only point that was out of it is the 673 average speed of the four-roots-per-axis mangrove at ebb tide (Fig.11 B). Even if the 674 seafloor current speed could be divided by three (0.63 to 0.23 m.s⁻¹), all the values 675 were really close on the diagram because of the logarithmic scale. 676

677

678 Fig.11 (small column size)



The value of a speed vector with two dimensions was calculated to combine results at 680 both tides to class the efficiency of the wave breakers. The vector's coordinates were 681 the average speed at flood tide and the average speed at ebb tide. The length of this 682 vector has been calculated with the following equation (Eq. 5). The lower the average 683 speed is, the better because it means a more important decrease in the swell erosion. 684 On the one hand, this method is interesting because even if a wave-breaker creates 685 an important slowdown at one tide, it will need to be performant at both tides to get the 686 lowest average speed possible. On the other hand, if it allows estimating the efficiency 687 of the model, this parameter doesn't give data about the increasing erosion of non-688 protected areas. 689

690

691
$$Eq.5: AS = ((ASFT)^2 + (ASET)^2)^{1/2}$$

692

693 With *AS* the average speed, *ASFT* the average speed at flood tide, and *ASET* the 694 average speed at ebb tide. The three parameters are in $m.s^{-1}$.

When the average speed for all the wave breakers was compared on a graph (Fig.12), it was very clear that the hybrid bio-wave-breaker is counter-productive and that the best set-up was two rows of the 4-roots-per-axis mangroves that had an average speed of 0.49 m.s⁻¹. The control value was 0.82 m.s⁻¹. Because of this and of the fact that this model didn't create important accelerations at any observation points, this set-up seemed to be the optimum one. Furthermore, it showed that the nest-inspired wave
 breakers didn't reduce a lot the average speed in addition to increasing the erosion in
 some non-protected areas.

703

Average speed per wave-breaker Hybrid bio-wave-breaker 0,86 Control 0,82 2-roots-per-axe mangrove 0,81 Quarters-of-nest 0,81 0,80 Elongated tubes Trapeze shaped 0,79 Straight bio-wave-breaker 0,77 Halves-of-nest 0,75 10-roots-per-axe mangrove 0,70 8-roots-per-axe mangrove 0,70 Chaotical mangrove 0,61 4-roots-per-axe mangrove 0,49 0,20 0,40 0,00 0,60 0,80 1,00 Average speed $(m.s^{-1})$

Fig.12 (small column size)

705

It could have been interesting to compare these seafloor current speed results with the height of waves when they break down and the distance to the wave-breaker. Collecting these values was not possible because of the meshing size in the simulation. The computer was able to calculate all the fluids values in every mesh, but it couldn't give The a wave picture we could measure precisely. Doing smaller simulations with a smaller meshing size could be a prospect of a future study.

712

713 4.5 Limits and prospects

714

It is important to notice that all these results were given by numerical simulations so there are a few limits to consider. Firstly, the mesh was 1 m large in all the simulations and 0.25 m large around the wave breakers. The results could have been different with a smaller meshing size. It was not possible to use a smaller one because the software would not have been able to generate the Lattice domain. Then, the length of the simulations was 45 seconds and they took between 10 and 19 hours to be computed. To study a global phenomenon such as erosion, definitive results couldn't be calculated and given without an experiment lasting several days.

Finally, the stability parameter was not optimum. This parameter in a numerical simulation depends on three conditions. If one is not satisfied somewhere in the generated domain, the parameter will reach 1.00 and higher. In the simulations, it was most of the time under 0.06 but sometimes, reach 0.2. The most important in a numerical simulation is to keep it between 0 and 1 but the lower this value is, the better.

Simulations' results need to be converging and to be verified experimentally to be 728 completely correct. The convergence is given by the software used (Xflow®2020). The 729 validation criteria was validated with the comparing of simulation results and 730 experimental results found in the scientific literature. The next step will be to validate 731 the results with test bench experiments. Another step could be to model a particular 732 georeferenced beach with details and to make simulations with precise swell and tide 733 data to obtain results about the possibility of the implementation of those strucures in 734 the reality. 735

Other prospects could be to use these bio-wave-breaker shapes in other software to 736 study the sedimentary outflow, to make larger-scale simulations with more than one 737 wave-breaker. It could be to make smaller and precise simulations to study the 738 behaviour of the current when it passes through a mangrove-shaped element, to make 739 740 simulations with a more complex beach shape. Another prospect could be to take one particular beach, to modelize it with details in 3D and to run simulations with precise 741 swell data to obtain results about a specific case. It has also been shown by many 742 authors that it is possible to use artificial reefs as wave breakers and to use wave 743 breakers as artificial reefs. However, in order to make artificial reefs specific materials 744 need to be used (Frihy et al., 2004; Düzbastılar. F. O. and Şentürk. U., 2009.; 745 Murakami and Maki, 2011). This article does not consider the material composition of 746 the wave breaker, therefore, the possibility of making artificial reef has not been 747 studied. 748

749 5. Conclusions

750

This paper aims to suggest bio-inspired wave breakers as a proof of concept in order 751 to reduce coastal setback and discuss about risks and opportunities of biomimicry 752 solutions. Biomimicry isn't a goal in itself but rather a potential solution for modern 753 754 problems, keeping in mind that all imitations are not useful. The results have shown the problems generated by existing models and especially the increasing erosion in 755 the non-protected areas. The Torquigener nest's imitation was a promising idea but 756 757 simulations have shown that these wave-breaker shapes accelerated the current in some areas. These results are not definitive and will need to be confirmed, but this 758 759 idea doesn't seem to bring a solution to the issue of the increasing erosion of the non-760 protected areas.

On the other side, the imitation of the mangroves showed interesting results. Both chaotical and symmetrical shapes caused an important general slowdown at ebb tide and an average slowdown at flood tide. The best set-up was two lines of a symmetrical model with four vertical roots by axis. This set-up could be perfected with a generative
design study and experimental data with a bench test to confirm simulations' results. It
would be interesting to make an economical study to see if the implementation of this
type of bio-wave breakers could be possible to solve the problem of coastal setback.

This paper built upon the idea that nature could give inspirations to solve natural problems. Even so, all the inspirations found in nature are not a way to deal with these problems. This study was set up in a mesotidal area. These results have been generated with general parameters that could feat with the Bay of Biscay in France. It could be different in the case of a precise location with detailed data. Furthermore, only the swash and the backwash have been studied. Drift currents have not been considered and could change the potential efficiency of these bio-wave breakers.

775 Credit author statement

776

Corentin Thomas: Conceptualization, Methodology, Software, Formal analysis,
 Investigation, Writing - Original Draft. Victor Lieunard: Investigation, Designing
 CATIA® Models, Writing - Review & Editing. Baptiste Oudon: Simulation, Writing –
 Review & Editing. Olivier Bain: Validation, Writing - Review & Editing. Rejanne Le
 Bivic: Validation, Writing - Review & Editing. Arnaud Coutu: Validation, Writing Review & Editing, Supervision, Project administration

783 Acknowledgements

784

The authors gratefully thank Dassault Systèmes Foundation for the financing, GéoLab of UniLaSalle Beauvais for the support provided for this work and all the students who previously worked on bio-inspired simulation subjects: J. M. Vient, C. Adam and L. Donovan, M. Arnaud and J. Michel.

789 References

790

Anthony, E.J., 2004. Sediment dynamics and morphological stability of estuarine

mangrove swamps in Sherbro Bay, West Africa. Marine Geology, Material Exchange

- 793 Between the Upper Continental Shelf and Mangrove Fringed Coasts with Special
- Reference to the N. Amazon-Guianas Coast 208, 207–224.
- 795 https://doi.org/10.1016/j.margeo.2004.04.009 .
- Aubie, S., & Tastet, J. P. (2000). Coastal erosion, processes and rates: an historical
 study of the Gironde coastline, southwestern France. Journal of Coastal Research,
 756-767
- Balouin, Y., Longueville, F., Colombet, Y., 2016. Video assessment of nearshore andbeach evolution following the deployment of a submerged geotextile wave breaker.
- 301 Journal of Coastal Research, Proceedings of the 14th International Coastal
- 802 Symposium (Sydney, Australia) Special Issue, 617–621. <u>https://doi.org/10.2112/SI75-</u>

803 <u>124.1</u>

Boulet, D., Giunta Fornasin, M.E., Heurtefeux, H., 2018. Modes de gestion et
responsabilités des ouvrages de lutte contre l'érosion et la submersion marine en
Occitanie. DOI : 10.5281/zenodo.2222442 .

Brivois, O., Idier, D., Thiébot, J., Castelle, B., Le Cozannet, G., & Calvete, D. (2012).
On the use of linear stability model to characterize the morphological behaviour of a
double bar system. Application to Truc Vert beach (France). *Comptes Rendus*

810 *Geoscience*, *344*(5), 277-287.

Butel, R., Dupuis, H., & Bonneton, P. (2002). Spatial variability of wave conditions on
the French Atlantic coast using in-situ data. *Journal of Coastal Research*, (36
(10036)), 96-108.

- Carpentier, L., Brosselard-Faidherbe, F., 1986. Influence de la période de la houle
 sur la stabilité des digues à talus. La Houille Blanche 371–375.
- 816 https://doi.org/10.1051/lhb/1986039.
- Castelle, B., Bonneton, P., 2006. Modelling of a rip current induced by waves over a
 ridge and runnel system on the Aquitanian Coast, France. C. R. Geoscience 338
 (10), 711–717.
- Cerc., 1984. Shore Protection Manual. Coastal Engineering Research Center, US
 Army Engineer Waterways Experiment Station, Vicksburg, MS.
- Charles, E., Idier, D., Thie bot, J., Le Cozannet, G., Pedreros, R., Ardhuin, F., Planton,
 S., 2011. Present wave climate in the Bay of Biscay: spatio- temporal variability and
- trends from 1958 to 2001. J. Climate, http:// dx.doi.org/10.1175/JCLI-D-11-00086.1.
- 825 Chen, S., Doolen, G.D., 1998. Lattice Boltzmann Method for Fluid Flows. Annual
- 826 Review of Fluid Mechanics 30, 329–364.
- 827 https://doi.org/10.1146/annurev.fluid.30.1.329
- Cho, K.-J., Wood, R., 2016. Biomimetic Robots, in: Springer Handbook of Robotics,
 Springer Handbooks. Springer International Publishing, Cham, pp. 543–574.
 https://doi.org/10.1007/978-3-319-32552-1_23
- Church, J.A;, P.U. Clark, A. Cazenaze, J.M. Gregory, S. Jevrejeva, A. Levermann,
- M.A. Merrifield, G.A. Milne, R.S. Nerem, P.D. Nunn, A.J. Payne, W.T. Pfeffer, D.
- 833 Stammer, and A.S. Unnikrishnan, 2013: Sea level change. Climate Change 2013:
- The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment
- 835 Report of the Intergovernmental Panel on Climate Change.
- Balrymple, Robert A., James T. Kirby, and Paul A. Hwang., 1984. Wave diffraction
 due to areas of energy dissipation. *Journal of waterway, port, coastal, and ocean*
- 838 *engineering* 110.1: 67-79.
- 839 Dassault Systèmes, 2021. Xflow 2021 User Guide.
- Delage, G., 1954. Using a breakwater for beach protection. Coastal Engineering
 Proceedings 35–35. https://doi.org/10.9753/icce.v5.35.
- Dodet, G., Bertin, X., Bouchette, F., Gravelle, M., Testut, L., Wöppelmann, G., 2019.
- 843 Characterization of Sea-level Variations Along the Metropolitan Coasts of France:

- Waves, Tides, Storm Surges and Long-term Changes. Journal of Coastal Research
 88, 10–24. https://doi.org/10.2112/SI88-003.1
- Bupuis, H., Michel, D., Sottolichio, A., 2006. Wave climate evolution in the bay of
 biscay over two decades. J. Mar. Syst. 63, 105–114.
- EMODnet, Mean depth full coverage, 2020 version, 2020, https://portal.emodnetbathymetry.eu/ (accessed 23 September 2021).
- Düzbastılar. F. O., Şentürk. U., 2009. Determining the weights of two types of
- artificial reefs required to resist wave action in different water depths and bottom
 slopes. Ocean Engineering, vol. 36, pp. 900–913.
- Earle, S., 2015. Physical Geology, 1st ed. BCcampus.
 https://opentextbc.ca/geology/ (accessed 18 January 2022).
- 855 Everett, T., Chen, Q., Karimpour, A., Twilley, R., 2019. Quantification of Swell Energy
- and Its Impact on Wetlands in a Deltaic Estuary. Estuaries and Coasts 42, 68–84.
 https://doi.org/10.1007/s12237-018-0454-z
- Frihy, O. E., El Ganaini, M. A., El Sayed. W. R., Iskander. M. M., 2004. The role of
 fringing coral reef in beach protection of Hurghada, Gulf of Suez, Red Sea of Egypt.
 Ecological Engineering, vol. 22, no. 1, pp. 17–25.
- Grondeau, M., Mercier, P., Guillou, S., Mear, Y., Poizot, E., n.d., 2016. Quelle
- turbulence ambiante pour la simulation numérique LBM-LES d'un environnement
 hydrolien? 11. http://website.ec-nantes.fr/actesjh/images/15JH/Articles/grondeau.pdf
 (accessed 22 September 2021)
- 864 (accessed 23 September 2021).
- Hay, C.C., Morrow, E., Kopp, R.E., Mitrovica, J.X., 2015. Probabilistic reanalysis of
- twentieth-century sea-level rise. Nature 517, 481–484.
- 867 https://doi.org/10.1038/nature14093
- Hegde, A.V., 2010. Coastal erosion and mitigation methods Global state of art.
 INDIAN J. MAR. SCI. 39.
- Heubes, D., Bartel, A., Ehrhardt, M., 2013. An introduction to Lattice-Boltzmann
- 871 methods. Novel Trends in Lattice-Boltzmann Methods Reactive Flow,
- 872 Physicochemical Transport and Fluid-Structure Interaction 3, 3–30.
- 873 https://doi.org/10.1016/j.cam.2013.09.019 .
- Hjulström, F., 1935. Studies of Morphological Activity of Rivers as Illustrated by the
 River Fyris. Bulletin of the Geological Institute University of Uppsala, 25, 221-527
- Hong Phuoc, V.L., Massel, S.R., 2006. Experiments on wave motion and suspended
- sediment concentration at Nang Hai, Can Gio mangrove forest, Southern Vietnam.Oceanologia 48.
- 879 https://www.researchgate.net/publication/26473444_Experiments_on_wave_motion_
- and_suspended_sediment_concentration_at_Nang_Hai_Can_Gio_mangrove_forest
- Southern_Vietnam (accessed 23 September 2021)
- 882 Hornack, M., 2011. Wave reflection characteristics of permeable and impermeable
- submerged trapezoidal breakwaters ProQuest (PhD thesis), Clemson University.

885

- Horstman, E.M., Dohmen-Janssen, C.M., Narra, P.M.F., van den Berg, N.J.F.,
- Siemerink, M., Hulscher, S.J.M.H., 2014. Wave attenuation in mangroves: A
- quantitative approach to field observations. Coastal Engineering 94, 47–62.
 https://doi.org/10.1016/j.coastaleng.2014.08.005.
- Isebe, D., Bouchette, F., MOHAMMADI, B., Azerad, P., Lambert, A., BUJAN, N.,
- Grasso, F., MICHALLET, H., 2008. Une nouvelle approche pour la protection des
- 892 plages : Application à la plage du Lido de Sète.
- 893 https://doi.org/10.5150/jngcgc.2008.025-l
- Jamshed, S., 2015. Introduction to CFD, in: Using HPC for Computational Fluid Dynamics. pp. 1–20. https://doi.org/10.1016/B978-0-12-801567-4.00001-5.
- Kawase, H., Mizuuchi, R., Shin, H., Kitajima, Y., Hosoda, K., Shimizu, M., Iwai, D.,
- Kondo, S., 2017. Discovery of an Earliest-Stage "Mystery Circle" and Development of
- the Structure Constructed by Pufferfish, Torquigener albomaculosus (Pisces:
- Tetraodontidae). Fishes 2, 14. https://doi.org/10.3390/fishes2030014 .
- 800 Kawase, H., Okata, Y., Ito, K., 2013. Role of Huge Geometric Circular Structures in
- the Reproduction of a Marine Pufferfish. Sci Rep 3, 2106.
- 902 https://doi.org/10.1038/srep02106.
- Komar P.D., Inman D.L., 1970. Longshore sand transport on beaches, J. Geophys.Res., 75(30), 5914-5927.
- Komar., 1979. Beach-slope dependence of longshore current, Journal of Waterway,
 Port, Coastal, and Ocean Division, Amer. Soc. Civil Engrs., 105(WW4), 460-464.
- Kopp, R.E., Horton, R.M., Little, C.M., Mitrovica, J.X., Oppenheimer, M., Rasmussen,
 D.J., Strauss, B.H., Tebaldi, C., 2014. Probabilistic 21st and 22nd century sea-level
 projections at a global network of tide-gauge sites. Earth's Future 2, 383–406.
 https://doi.org/10.1002/2014EF000239
- Lacroix, D., Mora, O., de Menthiere, N., Bethinger, A., 2019. La montée du niveau de
- la mer : conséquences et anticipations d'ici 2100, l'éclairage de la prospective.
- https://archimer.ifremer.fr/doc/00598/70975/69207.pdf (accessed 23 September
 2021).
- Larson, M., Kraus, N.C., 1992. Dynamics of Longshore Bars 2219–2232.
- 916 https://doi.org/10.1061/9780872629332.169
- 917 Lazure, P., Desmare, S., 2012. Caractéristiques et état écologique Golfe de
- 918 Gascogne-Courantologie.
- 919 https://dcsmm.milieumarinfrance.fr/content/download/4946/file/GDG_EE_06_Courant
- 920 ologie.pdf (accessed 23 September 2021).
- Le Cozannet, G., Lecacheux, S., Delvallée, E., Charles, E., Desramaut, N., Idier, D.,
- Krien, Y., Pedreros, R., Oliveros, C., 2011. Influence de la variabilité climatique sur
- les vagues dans le golfe de Gascogne 11.

- Le Cozannet, G., Lecacheux, S., Delvalle e, E., Desramaut, N., Oliveros, C., Pedreros, R., 2011. Teleconnection pattern influence in the Bay of Biscay. J. Climate,
- 926 doi:10.1175/2010JCLI3589.1.
- 927 Lee, W.K., Tay, S.H.X., Ooi, S.K., Friess, D.A., 2021. Potential short wave
- attenuation function of disturbed mangroves. Estuarine, Coastal and Shelf Science,
- Mangroves and People: Impacts and Interactions 248, 106747.
- 930 https://doi.org/10.1016/j.ecss.2020.106747 .
- Longuet-Higgins M.,1970. Longshore currents generated by obliquely incident sea
 waves, Journal of Geophysical Research, 75(33), 6778-6789.
- Masselink, G., Short, A.D., 1993. The effect of tide range on beach morphodynamics, a conceptual model. Jou*rnal of Coastal Research* 9, 785–800.
- Massel, S., Gourlay, M., 2000. On the modelling of wave breaking and set-up on coral reefs. Coastal Engineering 39. https://doi.org/10.1016/S0378-3839(99)00052-6
- 937 Matsuura, K., 2015. A new pufferfish of the genus Torquigener that builds "mystery
- 938 circles" on sandy bottoms in the Ryukyu Islands, Japan (Actinopterygii:
- 939 Tetraodontiformes: Tetraodontidae). Ichthyol Res 62, 207–212.
- 940 https://doi.org/10.1007/s10228-014-0428-5.
- 941 Mendez, Fernando J., and Inigo J. Losada., 2004. "An empirical model to estimate
- the propagation of random breaking and nonbreaking waves over vegetationfields." *Coastal Engineering* 51.2: 103-118.
- 944 Migniot, C., Lorin, J., 1979. Évolution du littoral de la côte des Landes et du Pays
- Basque au cours des dernières années. La Houille Blanche 65, 267–279.
 https://doi.org/10.1051/lhb/1979026.
- 947 Milieumarinfrance, 2020. Gestion intégrée du trait de côte.
- 948 https://www.milieumarinfrance.fr/Nos-rubriques/Actions-concretes/Gestion-integree-
- 949 du-trait-de-cote (accessed 23 September 2021).
- Miller, K.G., Kopp, R.E., Horton, B.P., Browning, J.V., Kemp, A.C., 2013. A geological
- 951 perspective on sea-level rise and its impacts along the U.S. mid-Atlantic coast.
- 952 Earth's Future 1, 3–18. https://doi.org/10.1002/2013EF000135
- Norris, Benjamin K., Julia, C., Karin, B., Henderson, S, M., 2021. Relating millimeterscale turbulence to meter-scale subtidal erosion and accretion across the fringe of a coastal mangrove forest. *Earth Surface Processes and Landforms* 46.3: 573-592.
- Pedreros, R., Howa, H., Michel, D., 1996. Application of grain-size-trend analysis for
 the determination of sediment transport pathways in intertidal areas. Mar. Geol. 135,
 35–49.
- 959 Phan, L.K., van Thiel de Vries, J.S.M., Stive, M.J.F., 2014. Coastal Mangrove
- 960 Squeeze in the Mekong Delta. Journal of Coastal Research 31, 233–243.
- 961 https://doi.org/10.2112/JCOASTRES-D-14-00049.1.
- Rasidah, K.W., Fakhri, I.M., Suhaimi, W.C., Jeyanny, V., Rozita, A., Fadzly, A.K.A.,
 2010. Soil physical changes of a coastal mudflat after wave breaker installation 4.

- Repousis, E., Metallinos, A., Memos, C., 2014. Wave Breaking over SubmergedBreakwaters.
- Sabatier, F., Samat, O., Brunel, C., Heurtefeux, H., Delanghe, D., 2009.
- 967 Determination of set-back lines on eroding coasts. Example of the beaches of the
- 968 Gulf of Lions (French Mediterranean Coast). Journal of Coastal Conservation 13, 57–
- 969 64. https://doi.org/10.1007/s11852-009-0062-y
- 970 Satjaritanun, P., Bringley, E., Regalbuto, J.R., Regalbuto, J.A., Register, J., Weidner,
- J.W., Khunatorn, Y., Shimpalee, S., 2018. Experimental and computational
- investigation of mixing with contra-rotating, baffle-free impellers. Chemical
- 973 Sénéchal, N., Gouriou, T., Castelle, B., Parisot, J. P., Capo, S., Bujan, S., & Howa,
- H. (2009). Morphodynamic response of a meso-to macro-tidal intermediate beach
 based on a long-term data set. *Geomorphology*, *107*(3-4), 263-274.
- 976 Stocker, T.F., D. Qin, G.-K. Plattner, M. 6 Tignor, S.K. Allen, J. Boschung, A. Nauels,
- Y. Xia, V. Bex, and P.M. Midgley, Eds. 7 Cambridge University Press, Cambridge,
 United Kingdom and New York, NY, USA, 1137– 8 1216.
- 979 http://www.climatechange2013.org/report/full-report/ (accessed 23 September
- 2021)Sweet, W., Horton, R., Kopp, R., Romanou, A., 2017. Sea level rise.
- 981 Publications, Agencies and Staff of the U.S. Department of Commerce.
- Thom, B.G., 1967. Mangrove Ecology and Deltaic Geomorphology: Tabasco, Mexico.
 Journal of Ecology 55, 301–343. https://doi.org/10.2307/2257879.
- Toyoshima, O., Shuto, N., Hashimoto, H., 1966. Wave Run-Up on Coastal
- 985 Structures. Coastal Engineering in Japan 9, 119–126.
- 986 https://doi.org/10.1080/05785634.1966.11924677.
- Van Rijn, L.C., 2011. Coastal erosion and control. Ocean & Coastal Management,
 Concepts and Science for Coastal Erosion Management (Conscience) 54, 867–887.
 https://doi.org/10.1016/j.ocecoaman.2011.05.004 .
- Van Santen, P., Augustinus, P.G.E.F., Janssen-Stelder, B.M., Quartel, S., Tri, N.H.,
- 2007. Sedimentation in an estuarine mangrove system. Journal of Asian Earth
- Sciences, Morphodynamics of the Red River Delta, Vietnam 29, 566–575.
- 993 https://doi.org/10.1016/j.jseaes.2006.05.011 .
- Vandermeirsch, F., Bonnat, A., Xiaoming, Y., Lazure, P., 2012. Caractéristiques et
 état écologique Golfe de Gascogne-Variation spatio-temporelle de la température et
- de la salinité. https://archimer.ifremer.fr/doc/00328/43924/43497.pdf (accessed 23
 September 2021).
- Wen, L., Weaver, J.C., Lauder, G.V., 2014. Biomimetic shark skin: design, fabrication
 and hydrodynamic function. Journal of Experimental Biology 217, 1656–1666.
- 1000 https://doi.org/10.1242/jeb.097097
- 1001 Williams, A.T., Rangel-Buitrago, N., Pranzini, E., Anfuso, G., 2018. The management
- of coastal erosion. Ocean & Coastal Management, SI: MSforCEP 156, 4–20.
- 1003 https://doi.org/10.1016/j.ocecoaman.2017.03.022.

- 1004 Wright, L. D., & Short, A. D. (1984). Morphodynamic variability of surf zones and 1005 beaches: a synthesis. *Marine geology*, *56*(1-4), 93-118.
- Yang, Judy Q., and Heidi M. Nepf., 2018. A turbulence-based bed-load transport
 model for bare and vegetated channels. *Geophysical Research Letters* 45.19: 10428.
- Zhang, Y., Yang, Y., Yang, K., Tan, X., Sun, X., Leng, B., Zhou, C., Zhu, B., 2020.
 Non-linear wave attenuation quantification model improves the estimation of wave
 attenuation efficiency of mangroves. *Estuarine, Coastal and Shelf Science* 245:
 106927.
- 1013 Figure Captions
- 1014 (Colors should be used for every figure)
- 1015 Table 1. Value of the wave breakers' parameters used in the simulations
- Table 2: Change of the mean water velocity on the probes for the 4 roots per axemangrove wave breakers
- 1018 Fig.1. 2D Profile of the domain and boundary conditions
- 1019 Fig.2. 2D Profile and 3D Model of the beach used in the simulations
- A. 2D Profile of the beach in the Golfe de Gascogne around ArcachonB. 3D Model of the beach with the location of the 8 observation points
- 1022 (probes) lines
- 1023 C. Location of the beach profile
- 1024 Fig.3. Inspired by the nest wave breakers
- 1025 A. Trapeze-shaped wave-breaker
- B. Elongated tubes wave-breaker
- 1027 C. Torquigener albomaculosus' nest
- D. First straight bio-inspired wave-breaker
- 1029 E. Second straight bio-inspired wave-breaker
- 1030 F. Wave-breaker built with quarters of the nest
- 1031 G. Wave-breaker built with thirds of the nest
- 1032 H. Wave-breaker built with halves of the nest
- 1033 I. Wave-breaker with the taller dune facing the sea
- 1034 J. Hybrid wave-breaker
- 1035 Fig.4. Inspired by the mangrove wave breakers
- 1036A. Mangrove scheme
- B. Symmetrical mangrove with eight vertical roots by axis
- 1038 C. Symmetrical mangrove with ten vertical roots by axis
- D. Symmetrical mangrove with four vertical roots by axis
- 1040 E. Symmetrical mangrove with two vertical roots by axis
- 1041 F. Chaotical mangrove
- 1042 Fig.5. Seafloor current speed depending on the existing wave-breaker used

1043	Blue: control seafloor current speed						
1044	Purple: seafloor current speed with the trapeze-shaped wave-breaker						
1045	Yellow: seafloor current speed with the elongated tubes wave-breaker						
1046 1047	A. Results at flood tideB. Results at ebb tide						
1048	Fig.6. Seafloor current speed depending on bio-wave-breaker used						
1049	Blue: control seafloor current speed						
1050	Red: seafloor current speed with the straight bio-wave-breaker (Fig.2 D)						
1051 1052	Purple: seafloor current speed with the quarter-of-nest bio-wave-breaker (Fig.2 F)						
1053	Green: seafloor current speed with halves-of-nest bio-wave-breaker (Fig.2 H)						
1054	Yellow: seafloor current speed with hybrid wave-breaker (Fig.2 J)						
1055	A. Results at flood tide						
1056	B. Results at ebb tide						
1057	Fig.7. Seafloor current speed depending on the roots' density						
1058	Blue: control seafloor current speed						
1059	Purple: seafloor current speed with four roots per axis						
1060	Yellow: seafloor current speed with eight roots per axis						
1061	Green: seafloor current speed with ten roots per axis						
1062	A. Results at flood tide						
1063	B. Results at ebb tide						
1064	Fig.8. Seafloor current speed depending on the number of lines						
1065	Blue: control seafloor current speed						
1066	Purple: seafloor current speed with one line						
1067	Yellow: seafloor current speed with two lines						
1068	Green: seafloor current speed with three lines						
1069 1070	A. Results at flood tideB. Results at ebb tide						
1071 1072	Fig.9. Variation of seafloor current velocity per probe.4 roots per axe mangrove wave breaker simulation with different current orientation						
1073	Blue: Seafloor current speed for the 75° simulation						
1074	Orange: Seafloor current speed for the 60° simulation						

- 1075 Gray: Seafloor current speed for the 45° simulation
- 1076 Yellow: Seafloor current speed for the 30° simulation
- 1077 Fig.10. Average speeds per observation point (probe)with and without two rows of 1078 two-vertical-roots-per-axis mangroves
- 1079 Green: average speed per point in the control simulation
- 1080 Orange: average speed per point in the simulation with wave-breaker
- Blue: average acceleration when the wave-breaker is used
- A. Results at flood tide
- B. Results at ebb tide
- Fig.11. Simulations' average speed in a Hjulström-Sundborg diagram (Earle, 2015;
 Hjulström, 1935)
- 1086 A: All the simulations except one
- B: Simulation at ebb tide of a 4-roots-per-axis mangroves wave-breaker
- 1088 Fig.12. Average speed per wave-breaker
- 1089 Green: under 0.65 m.s⁻¹
- 1090 Yellow: between 0.66 and 0.75 m.s⁻¹
- 1091 Orange: between 0.76 and 0.85 m.s⁻¹
- 1092 Red: over 0.86 m.s⁻¹