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1. Introduction

Trace metal and metalloid elements (TMM) are naturally present in rivers as a result of mechanical erosion processes or chemical alteration of uncontaminated rocks and soils. They also originate from multiple sources related to anthropogenic activities resulting in the contamination of rivers to varying degrees of severity (Salomons and Forstner, 1984). Anthropogenic impacts are sometimes very old and may date back to the Roman era (Elbaz-

Poulichet et al., 2011), the Bronze Age (Elbaz-Poulichet et al., 2020), or even prehistory (Chiaradia et al., 2003). However, many studies agree that the impacts from the last 150 years, since the industrial revolution of western countries, have been much more significant (Foster and Charlesworth, 1996). Releases of TMM can be in dissolved and/or particulate form. Particulate discharges may be transported to the lower reaches of a watershed depending on hydraulic conditions, leading to the storage of these particles in accumulation areas within rivers, such as reservoirs behind dams, floodplains or ponds (Foster and Charlesworth, 1996). Using stored sediments, the reconstruction of temporal trends of metal contamination within a watershed is possible when the latter are undisturbed. Stored sediments with high accumulation rates (e.g., $\geq$1 cm y^{-1}) are only slightly affected by the early diagenesis process (Callender, 2000), thus these sediments can be used to trace temporal trends of metal contamination. This has been done for several large French rivers (e.g., Garonne River: Grousset et al., 1999; Audry et al., 2004; Seine River: Le Cloarec et al., 2011; Ayrault et al., 2012; Loire River: Grosbois et al., 2012; Dhivert et al., 2016).

The Eure River in Normandie, is a major tributary of the Seine estuary and is part of the Seine River watershed. This watershed is known for a history of significant metal contamination within its fluvial (Le Cloarec et al., 2011) and estuarine parts (Chiffoleau et al., 2012; Vrel, 2012).

The objective of this study is to assess and characterise the anthropogenic impacts within the Eure River watershed over the last 80 years. The quantification of current inputs from the Eure watershed into the Seine estuary was determined through TMM concentrations on suspended particulate matter (SPM), to enable determination of particulate metal flux. The reconstruction of TMM temporal trends was undertaken using sediment cores collected in the lower reaches of the watershed. The determination of natural values of TMM within the watershed allowed the assessment of current and past anthropogenic impacts. The magnitude

of the impact could also be traced over time based on the estimation of past theoretical flux. The sources of the recorded contamination were determined by reconstructing the industrial history of the watershed and using Pb isotope ratios.

2. Material and Methods

2.1. Characteristics of the Eure watershed

The Eure River is a tributary of the Seine estuary, regulated upstream by the Poses Dam (Figure 1A). The Eure River watershed (6,017 km²) is the fourth largest tributary of the Seine River watershed in terms of surface area. It falls behind the Yonne, Marne and Oise watersheds but is the first contributing watershed in the estuarine section (i.e., the last 150 km before the mouth to the Channel). It is based on a chalk formation of the Upper Cretaceous, covered by Tertiary sandy-clay deposits and more recent deposits dating back to the Quaternary period (Quesnel, 1997; Laignel et al., 1998). The Eure River waters are characterised by a conductivity of 568 ± 52 µS cm^{-1}, a pH equal to 8.1 ± 0.2 and a mean temperature of 13 ± 5 °C on the period from 1971 to 2019 (www.naiades.eaufrance.fr). Moreover, the Eure River has a mean annual flow (Q_{mean}) of 22.13 ± 6.69 m^3 s^{-1} based on the period from 1971 to 2019 (www.hydro.eaufrance.fr). It is the fifth largest tributary of the Seine watershed in terms of flow and its two major tributaries are the Avre and Iton rivers (Figure 1B).

2.2. Field Sampling Methods

2.2.1. Description of core sites and sampling methodology

Les Damps Pond is upstream of the Eure River spillway (Figure 1C). The pond, with an average depth of 50 cm and spanning an area less than 1 ha, continually accumulates SPM from the Eure watershed (Figure 1D). As the site is 2 km upstream of a spillway (vertical height: ~3 m), the tides have no noticeable effect, even in the case of a high tidal coefficient.

The Martot Pond, spanning an area of ~7 ha, is located in the lower reaches of the study area (Figure 1C). Unlike Les Damps Pond, this pond is affected by the Seine estuary during tidal flows through the Eure River outlet. However, the impact of the tide is limited and temporally variable, as its flow persists for about 3 h and does not occur every day. In addition, the Martot Dam, located 200 m downstream of the Martot Pond, prevented tidal flows for tidal coefficients lower than 70, until its removal in October 2017 (Figure 1E). Thus, to ensure Martot Pond was not impacted by the Seine River, results from this site were compared with those from Les Damps Pond. The Martot Dam also regulated the water level in the river, and the river was continuously connected to the Martot Pond, even during periods of low flow.

The La Forge Pond, located upstream of the Eure watershed, is crossed by the Avre River and is less than 10 km downstream of the headwaters of the Avre River (Figure 1F). This pond was constructed in anticipation of iron and steel activities during the seventeenth century. It is at the outlet of a watershed spanning less than 14 km² where the geology is similar to that of the entire watershed, i.e., an Upper Cretaceous formation composed of flint clay (Turonian) and sand (Cenomanian) outcropping near the river.

Sediment cores were extracted using a UWITEC® gravity corer and PVC tubes with a diameter of 90 mm (e.g., Sabatier et al., 2014). The water-sediment interface was preserved during coring. Coring was carried out in the following dates and locations:

- January 2015 (DAM15-02 core) and January 2017 (DAM17-02 core) for Les Damps Pond;

- January 2015 (MAR15-01 core) and February 2016 (MAR16-02 core) for Martot Pond; and
- June 2017 (FOR17-02 core) for La Forge Pond.

The characteristics of sediment cores are listed in Table 1.

2.2.2. Measurement of flow and turbidity

Streamflow of the Eure River (Q_{water}; $m^3\ s^{-1}$) was directly measured using an Argonaut-SL3000 current meter (Side-Looking Doppler Current Meters; SonTeck®) installed on the left bank of the Eure River at the Les Damps site. Measurements were taken at 15 min intervals from 27 January 2017 to 22 May 2018 (Figure 1D).

Turbidity (NTU) was measured using a SAMBAT multi-parameter probe (with self-cleaning sensors; NKE Instrumentation®) installed in the channel 90 cm below the surface at the Les Damps site. Measurements were taken at 15 min intervals during the same period as that of streamflow (Figure 1D).

2.2.3. SPM sampling

SPM were collected using an home-made time-integrated mass-flux sampler (TIMS), the design of which was based on Phillips et al. (2000) and Russell et al. (2000) (Figure S1 in supplementary materials). The central tube (1 m long, diameter of 10 cm) and the two outer inlet and outlet pieces are made of PVC. The pipe has a diameter of 4 mm, allowing water to enter and exit, is made of nylon. The assembly of the parts and waterproofing of the central tube was undertaken using O-rings without glue. The TIMS were installed on a raft in the channel 90 cm below the surface at the Les Damps site, at the same depth as the turbidity sensor (Figure 1D). All TIMS changeover dates are shown in italics in Figure 2A.

2.3. Analytical Methods and data processing

126

127 *2.3.1. Grain size analysis*

128 Grain size distribution can influence the behaviour of chemical substances. Short-lived
129 radionuclides (^{210}Pb, ^{137}Cs, and ^{241}Am) are preferentially adsorbed to fine particles (He and
130 Walling, 1996) as metallic and metalloid elements (Salomons and Forstner, 1984; Horowitz
131 and Elrick, 1987; Brügmann, 1995). Thus, grain size distribution can have a considerable
132 impact on chronology and TMM concentrations in sediment cores (He and Walling, 1996).

133 The grain size distribution of sediment cores was measured by laser diffraction (LS 13320
134 Particle Size Analyser Beckman Coulter™) across a range 0.04–2000 µm. Measurements
135 occurred every cm by integrating 1 cm on the DAM15-02 and MAR15-01 cores.

136

137 *2.3.2. Total organic carbon analysis*

138 Total organic carbon (TOC) and the extent of organic matter (OM) degradation can
139 influence TMM concentrations as some TMM have a strong affinity to TOC (Tseng et al.,
140 2001; Audry et al., 2006; Masson et al., 2011; Petit et al., 2013; Coynel et al., 2016). As such,
141 TOC contents are needed to determine whether variations in TMM in the sediment cores are
142 related to anthropogenic activities.

143 Sediment Cores were analysed using Rock-Eval 6 (RE6) pyrolysis at the ISTO laboratory,
144 University of Orléans ("Turbo" model RE6 pyrolyzer, Vinci Technologies), as per the method
145 described in Copard et al. (2006). TOC and the OM quality were measured on the DAM17-02
146 and MAR15-01 cores at intervals of 2 and 1 cm, respectively.

147

148 *2.3.3. XRF core scanning (XRF-CS)*

149 X-Ray fluorescence spectrometry (XRF) was used to semi-quantitatively determine the major
150 and trace element variations of sediment deposits (Richter et al., 2006). The XRF core

scanning data can be used to describe particulate TMM contamination in sediment cores (Lepland et al., 2010; Hennekam et al., 2019; Elbaz-Poulichet et al., 2020). The data presented were normalized by Ti to remove possible variations due to detrital inputs (Duan et al., 2014; Bábek et al., 2015). Ti was considered a better "normalizing element" than Al, which is used in coagulation/flocculation processes in wastewater treatment plants.

The relative contents of TMM were measured using an Avaatech XRF core scanner (EDYTEM, University of Savoie Mont Blanc). Measurements were taken every 5 mm for the MAR15-01 and DAM17-02 cores and every 2 mm for the MAR16-02 core.

2.3.4. Particulate trace metal and metalloid analysis

Microwave-assisted sediment digestion for particulate TMM analyses (for As, Cd, Cu, Cr, Ni, Mn, Pb, Zn), was performed on samples collected from the MAR16-02 and DAM17-02 cores. This consisted of 0.5 g of dry, powered and homogenised sediment with aqua regia (HNO_3: 3 mL TraceMetal™ Grade; HCl: 9 mL TraceMetal™ Grade, Fisher Scientific®). Following pseudo-total mineralization (residual fractions containing silica particles being only partially digested), solutions were diluted in 100 mL Milli-Q® water. Particulate TMM concentrations were measured using inductively coupled plasma-atomic emission spectroscopy (ICP-AES, iCAP 6000 Series, Thermo Fischer®) with external calibration. The accuracy and precision of measurements were verified with a certified reference material (Trace Element on Fresh Water Sediment CNS301-04-050, Sigma Aldrich) and results showed recoveries of ≥88% with a precision of ~3.1%.

Total digestions, including residual fractions, were performed on selected samples, collected at different depths for the DAM17-02 core, corresponding to the highest and lowest levels determined by ICP-AES. They were also conducted on sediment collected at the bottom of the FOR17-02 core. Representative samples consisting of 30 mg of dry, powdered and

176 homogenized sediment, were digested in closed polypropylene (PP) reactors (DigiTUBEs®,
177 SCP SCIENCE) in a heating block (2 h at 115 °C) with 1.5 ml HCl (12 *M* Suprapur®,
178 Merck), 0.75 ml HNO_3 (14 *M* Suprapur®, Merck) and 2.5 ml HF (22 *M* Suprapur®, Merck),
179 as described in previous studies (e.g., Coynel et al., 2016). Particulate TMM concentrations
180 were measured by inductively coupled plasma-mass spectroscopy (ICP-MS, X7 Series 2,
181 THERMO) with external calibration under standard conditions. The accuracy and precision of
182 measurements were verified with certified reference materials (NCS DC 2702, CRM 8704,
183 BCR 667) and results showed recoveries of ≥93% with a precision of ≤4% (relatively
184 standard deviation; RSD). The concentrations (for Cr, Ni, Cu, Zn, As, Cd, and Pb) determined
185 by ICP-MS (total concentrations) were always consistent with those determined by ICP-AES
186 (pseudo-total concentrations) being 1.1 to 1.5 times higher than the latter.
187 Sediment samples from MAR15-01 (n=6) were analysed for Pb radiogenic isotopes.
188 Reagents used for digestion and chemical procedures were prepared from sub-boiled distilled
189 concentrated hydrochloric and nitric acids using a Cleanacids® device (Analab, France)
190 except $HClO_4$ (Suprapur®, Merck) Approximately 100 mg of sample was weighed and
191 dissolved in savillex® Teflon beakers in a mixture of HF (24 *M*, Cleanacids®, Analab),
192 HNO_3 (14 *M*, Cleanacids®, Analab), and $HClO_4$ (12 *M*, Suprapur®, Merck) for four days at
193 160 °C on a hot plate. After evaporation to dryness, samples were dissolved in aqua regia and
194 heated for 24 h at 130 °C. Pb fractions were chemically separated following conventional
195 column chemistry procedures described in Gale (1996). Pb isotopic compositions analyses
196 was conducted at the Pôle de Spectrométrie Océan (PSO) in Brest, France using a multi-
197 collector ICP-MS (MC-ICPMS, Neptune, Thermo Scientific). Pb isotope ratios were
198 corrected for instrumental mass fractionation and machine bias using the Tl doping method of
199 White et al. (2000) and SRM981 Pb standard bracketing for every three samples. Pb isotope
200 reproducibility, based on 21 replicate analyses of NIST (National Institute of Standards and

Technology) SRM981 was 0.0018 (2σ) for $^{206}Pb/^{204}Pb$ and $^{207}Pb/^{204}Pb$ and 0.0045 (2σ) for $^{208}Pb/^{204}Pb$.

2.3.5. Suspended particulate matter and particulate TMM flux determination

To convert turbidity measurements into SPM concentrations (C_{MES}; mg L^{-1}), an NTU-SPM calibration line was established from SPM samples collected monthly at the Les Damps site (Figure S2). Equation (1) of this calibration line ($R^2 = 0.85$) is:

$$C_{MES} = 1.155 \times Turbidity \qquad \text{Equation (1)}$$

The monthly SPM flux ($F_{SPM;}$ t $month^{-1}$) was then determined according to Equation (2):

$$F_{SPM} = \sum Q_{water} \times C_{MES} \qquad \text{Equation (2)}$$

Particulate TMM flux was estimated monthly over the sampling period based on the SPM flux and TMM concentrations determined from the collected SPM.

2.3.6. Enrichment factors calculation

To assess the impact of anthropogenic activities based on concentrations in environmental samples, these concentrations need to be standardised to reduce the influence of the sedimentary matrix (Ackermann, 1980; Daskalakis and O'Connor, 1995). Normalisation is based on a conservative trace element (i.e., not influenced by oxidation-reduction or diagenesis processes in sediments). Variations are characterised in the fine fraction varying proportionally with background concentrations, where most TMM are preferentially enriched. This normalising element should not have anthropogenic inputs that alter its distribution. The element Th is highly insoluble (Van Calsteren and Thomas, 2006), and is known to have a strong negative correlation with particle size, which allows its effects to be corrected (e.g., Coynel et al., 2007; Larrose et al., 2010; Lanceleur et al., 2011a, 2011b; Le Cloarec et al., 2011). In the literature, enrichment factors (EFs) are commonly used to estimate the impact

of anthropogenic activities on the natural cycle of trace elements. EFs were determined according to Equation (3):

$$EF = \frac{TE \,/\, RE}{TE_{ref.background} \,/\, Th_{ref.\ background}} \qquad \text{Equation (3)}$$

TE and *RE* represent the concentration of TMM studied and the reference element (Th) chosen for a given sample. $TE_{ref.\ background}$ and $RE_{ref.\ background}$ represent the natural content of the studied element and the natural content of Th, respectively.

2.4. Core dating

2.4.1. Short-lived radionuclides

The dating of recent deposits within sediment cores and the estimation of sedimentation rates was carried out using radioelements with short periods. The sedimentation rate was determined from the activity of ^{210}Pb in excess ($^{210}Pb_{ex}$), by subtracting ^{226}Ra from the $^{210}Pb_{total}$ (Goldberg, 1963). Artificial radioelements ^{137}Cs and ^{241}Am were used to identify nuclear weapons tests to confirm the age model established from the $^{210}Pb_{ex}$ results. A peak was observed in 1963 AD and in 1986 AD when the Chernobyl nuclear power plant accident occurred (Appleby et al., 1991).

The activities of $^{210}Pb_{ex}$, ^{137}Cs and ^{241}Am were determined using gamma spectrometry with a germanium spectrometer in ultra-low background noise settings at the Laboratoire Souterrain de Modane (LSM). The age model was computed with *serac* R package (https://github.com/rosalieb/serac, (Bruel and Sabatier, 2020)). The MAR15-01 core was sampled from 0 to 80 cm at intervals of 2, 4, and 6 cm.

2.4.2. ^{14}C dating

250 ^{14}C dating was conducted at the Poznań Radiocarbon Laboratory from organic macro
251 remains collected at the bottom (i.e., 68–69.5 cm interval) of the FOR17-02 core. The
252 IntCal13 calibration curve (Reimer et al., 2013) was used for ^{14}C age calibration, and the
253 sample was determined to be from the period between 1646 and 1669 cal AD (Figure S3).
254
255 **3. Results and interpretation**
256
257 3.1. Determination of local geochemical background
258
259 The beginning of the FOR17-02 core (68–69.5 cm) was radiocarbon dated between 1646
260 and 1949 (1 sigma) with a maximum probability that it is between 1646–1669 (56%) and
261 1781–1798 (39%) (Figure S3). The 68–69.5 cm fraction was characterised by a median grain
262 (D_{50}) of 20.9 μm and a TOC content of 2.77% (dry weight; dw) (additional description of the
263 FOR17-02 core is provided in Figure S4).
264 The TMM concentrations of this sample are summarized in Table 2. In light of the small
265 geological variations within the Eure watershed and the sample being dated back to the pre-
266 industrial revolution period, the TMM concentrations are indicative of the geochemical
267 background (GB) of the Eure watershed. The choice of the GB in this study, differs from
268 some studies that have used levels measured at the bottom of sediment cores, considered
269 representative of the GB (e.g., Gascón Díez et al., 2017), or used a regional geochemical
270 baseline, representing levels measured prior to significant additional anthropogenic inputs, but
271 not corresponding with a regional GB (Larrose et al., 2010). In other cases, for example for
272 the Seine watershed (67,500 km²; Le Cloarec et al., 2011), the GB was the combination of
273 several average levels of TMM in fine sediments collected in forest streams draining a single
274 rock type (Horowitz et al., 1999; Thévenot et al., 2007).

275

276 3.2. Current behaviour of the Eure River

277

278 *3.2.1. Hydro-sedimentary behaviour*

279 The average flow in the Eure River was 22.0 $m^3 s^{-1}$ at the Les Damps site in 2017. There
280 were no major flood events during this year, with the exception of the month of March when
281 streamflow reached 86.7 $m^3 s^{-1}$, the highest in 2017. The beginning of 2018 was marked by a
282 50-year flood during which five flood peaks were recorded. Maximum flow values occurred
283 between late January and early February 2018, with streamflows close to 180 $m^3 s^{-1}$ (Figure
284 2A).

285 The monthly SPM flux also followed the same trend as streamflow. The maximum monthly
286 flux in 2017 was 4.27×10^4 t in March 2017, while the minimum flux was 812 t in July 2017.
287 The total flux for 2017 was 10.8×10^4 t, with a specific sediment yield (SSY) of about 18 t
288 $km^{-2} y^{-1}$. The 2018 flood (January–March) transported 36.0×10^4 t of SPM, of which more
289 than half (20.4×10^4 t) passed through in January (Figure 2A).

290

291 *3.2.2. Evolution of TMM concentrations*

292 Figure 2B presents the changes in TMM and Mn concentrations over the entire sampling
293 period at the Les Damps site. According to the correlation matrix, most of the TMM and Mn
294 did not show any correlation with Q_{mean}, SPM_{mean} and dw (for SPM collected by TIMS),
295 indicating that their concentration variations are not directly related to the hydro-sedimentary
296 behaviour of the Eure River, nor to the mass of SPM collected (Figure S5). Only As showed
297 positive but low correlations ($0.36 < R < 0.56$) with these three parameters. Overall, TMM
298 and Mn were positively correlated with each other ($0.42 < R < 0.99$), but some elements such
299 as V, As and Th showed particularly strong positive correlations ($0.90 < R < 0.99$) and

300 therefore co-variated during the sampling period. This was particularly the case during the 50-
301 year flood from 22 January to 19 February, 2018. During this time, only V, As and Th
302 reached maximum levels at 55.3, 8.45 and 6.44 mg kg^{-1}, respectively (Figure 2B). Chromium,
303 Co, Ni and Mn had strong correlations between them ($0.60 < R < 0.89$) and with Th or V for
304 Cr, Co and Ni ($0.73 < R < 0.88$). Lead was correlated with Ag ($R = 0.89$) and Sb ($R = 0.70$),
305 and like Ag, was negatively correlated with Q_{mean}, SPM_{mean} and dw. This resulted in low
306 levels of Pb and Ag during the flood with Ag having reached its lowest concentration of 0.75
307 mg kg^{-1} (Figure 2B).
308 The average concentrations of V, As and Th based on the SPM were close to those of the
309 GB of the Eure watershed (Table 2). The other TMM had average concentrations higher than
310 those of the GB. Over the entire sampling period, Cu, Cd and Pb measured at levels 12, 18,
311 and 14 times higher than the GB, respectively.
312 For concentrations of TMM normalised by Th, the average ratios were found to be higher
313 than those calculated for the GB, with the exception of V, As and Th. This indicates that the
314 inputs of As within the watershed were not of anthropogenic origin. For the other TMM, it is
315 likely that one or more additional anthropogenic sources releasing TMM into the river or that
316 old contaminated and stored/trapped sediments/soils have been remobilised. This may be
317 likely to occur during flood or work periods, and transit through the Eure River.
318
319 *3.2.3. Particulate TMM flux*
320 The total monthly flux of particulate TMM are listed in Table 3. In 2017, 50% of TMM flux
321 was transported between January and May as per the water flux. It may also have been
322 transported between January and April for V, As, Ag, Cd, Sb and Th as per the SPM flux.
323 Maximum flux occurred in March during the episodic flood. The flood in 2018 had a major
324 impact on the transport of TMM within the watershed. Indeed, between January and March,

325 the TMM flux was 3.3 times greater on average compared to the whole of 2017. Note that this
326 average ratio was 1.9 times if only the month of January is considered. Specifically, this
327 behaviour was more pronounced for V, Co, Ni, Cu, As, Th (ratio of 3.5–4.0) and less
328 pronounced for Ag (ratio of 2.0) and Pb (ratio of 2.8). During the flood event, more than half
329 (57 ± 1%) of the flux occurred in January, including the particulate flux for all the TMM. The
330 TMM with the highest flux was Zn and Pb, followed by Cr and Cu (Table 3).
331 While variations in TMM contents are independent of the hydro-sedimentary behaviour of
332 the Eure River, variations in TMM flux, such as increases during flood periods, are due to the
333 increase in hydro-sedimentary inputs. These streamflow increases are thus indicative of the
334 increased erosion of soils or sediments stored in the river.
335
336 3.3. Past contamination in the Eure River watershed
337
338 *3.3.1. Sedimentary characteristics*
339 The sediment deposits of the DAM15-02 core (Les Damps Pond) had a grain size
340 distribution mainly centred around D_{50}, equal to 29.5 ± 6.7 µm, with a constant TOC content
341 of 5.36 ± 0.47% (Figure 3A). The sediment deposits of the MAR15-01 core (Martot Pond),
342 were characterised by two sedimentary facies called U2 (77–138 cm) and U1 (0–77 cm)
343 (Figure 3B). Along U2, the grain size distribution was centred around a low D_{50} equal to 13.8
344 ± 7.0 µm, and there was a low and constant mean TOC content (1.75 ± 0.33%). Along U1, the
345 grain size distribution was also centred around a D_{50} of 33.9 ± 7.8 µm, and the mean TOC
346 content (3.88 ± 1.49%) appeared to increase slightly within U1 towards the interface. There
347 were particularly high values close to the interface (TOC = 8.26 ± 0.19% between 9 and 12
348 cm) (Figure 3B). The D_{50} and mean TOC content were higher in the Martot Pond for U1
349 compared to U2. However, the D_{50} of U1 in the Martot Pond and that of Les Damps Pond

were comparable, while the mean TOC content was higher in Les Damps Pond (Gardes et al., 2020a).

3.3.2. Age model

The activity of $^{210}Pb_{ex}$ was measured up to 80 cm (U1 of MAR15-01 core) and was found to decrease almost continuously with depth (Figure 4A). A sedimentation rate of 12.73 ± 2.2 mm y^{-1} (R^2 = 0.60) was estimated using the 'constant flux, constant sedimentation rate' (CFCS) model (Krishnaswamy et al., 1971). The ^{137}Cs and ^{241}Am activities peaked at 60 and 58 cm, respectively, corresponding to the maximum fallout from atmospheric nuclear weapon testing in 1963 (Robbins and Edgington, 1975). There was also a peak in the ^{137}Cs activity at 40 cm, potentially related to the Chernobyl accident in 1986 (Appleby et al., 1991) (Figure 4B). The first identification of ^{137}Cs activity dated 1955 (Pennington et al., 1973) corresponded to a depth of 71 cm. From the mean sedimentation rate obtained by the $^{210}Pb_{ex}$, an age model was constructed that was consistent with the artificial radionuclide data (Figure 4C). The transition between U1 and U2 was dated at 1954 ± 10 y. However, the history of the study site suggests that around 1939, the Martot Pond, then connected to the Seine River, was disconnected from the Seine River and subsequently connected to the Eure River in the following years. Thus, the U2 facies corresponds to sedimentary deposits from the Seine watershed while the U1 facies corresponds to contributions from the Eure watershed since the 1940s (Gardes et al., 2020a).

As the MAR16-02 and DAM17-02 cores on which the TMM were measured (see Section 3.3.3) were not dated, an age model was defined using the MAR15-01 core. This was done by linking the peaks of the Pb/Ti and Zn/Ti ratios together based on analogy, or using the U1/U2 transition level also visible along the MAR16-02 core (Figure S6). Among the several parameters useful to correlate sediment cores, XRF data is interesting because of the

375 measurements resolution. In addition, Pb and Zn already showed distinct temporal trends in
376 the Martot and Les Damps Ponds, using XRF analysed. Indeed, high levels of Zn were
377 obtained in the 1950s-1960s whereas high levels of Pb were obtained in the 1990s,
378 corresponding to the "middle" of the sediment cores and approximatively to the top of the
379 sediment cores, respectively (Gardes et al., 2020a). The results indicated average
380 sedimentation rates of 11.3 mm y^{-1} for the MAR16-02 core and 12.2 mm y^{-1} for the DAM17-
381 02 core.
382
383 *3.3.3. Evolution of the TMM concentrations along sediment cores*
384 • MAR16-02 core
385 The TMM concentrations in the MAR16-02 core were measured for U1 (0–77 cm),
386 corresponding to inputs from the Eure River into the Martot Pond, and for the last length of
387 U2 (77–90 cm), corresponding to inputs from the Seine River. Chromium, Mn, Ni, Cu, Zn
388 and Cd concentrations increased from the transition U2/U1 (Seine Unit/Eure Unit) with a
389 peak in the late 1950s (Cr = 211 ± 1 mg $kg^{-1;}$ Mn = 2198 ± 9 mg $kg^{-1;}$ Ni = 48.9 ± 0.1 mg $kg^{-1;}$
390 Cu = 399 ± 4 mg $kg^{-1;}$ Zn = 905 ± 1 mg $kg^{-1;}$ Cd = 42.3 ± 0.1 mg kg^{-1}), and in the mid-1950s
391 for Mn. Levels of Cr, Mn, Ni, Zn, and Cd remained high until the early 1960s and then
392 declined, in stages, until the late 1980s. From the 1980s onwards, the contents of Cr, Mn, Ni,
393 Cu, Zn and Cd remained stable (Cr = 56.9 ± 0.3 mg $kg^{-1;}$ Mn = 234 ± 1 mg $kg^{-1;}$ Ni = 22.6 ±
394 0.1 mg $kg^{-1;}$ Cu = 94.2 ± 0.5 mg kg^{-1}; Zn = 311 ± 1 mg $kg^{-1;}$ Cd = 3.78 ± 0.03 mg kg^{-1}) (Figure
395 5A). Arsenic concentrations had a peak in the U2 facies and then decreased during the
396 transition from the Seine Unit to Eure Unit with a peak in the late 1950s of 12.2 ± 0.2 mg kg^{-1}.
397 Concentrations of As then decreased until the end of the 1980s, stabilising around a mean
398 content of 4.12 ± 0.23 mg kg^{-1} up to the surface (Figure 5A). Pb concentrations increased
399 significantly from the late 1980s with a peak of 669 ± 3 mg kg^{-1} in the mid-2000s. During the

2010s the Pb concentrations declined rapidly but remained elevated and above levels measured prior to the increase in the late 1980s (Figure 5A).

- DAM17-02 core

The evolution of TMM levels in the Les Damps Pond are similar to those obtained for the Martot Pond, indicating a global signature of the watershed within these two ponds, not impacted by local inputs (e.g., around the ponds). Indeed, the concentrations of Cr, Mn, Ni, Cu, Zn, Cd were similar with the maxima during the 1960s and 1970s (Cr = 171 ± 1 mg kg^{-1}; Mn = 1283 ± 2 mg kg^{-1}; Ni = 75.4 ± 0.2 mg kg^{-1}; Cu = 490 ± 1 mg kg^{-1}; Zn = 992 ± 3 mg kg^{-1}; Cd = 54.6 ± 0.1 mg kg^{-1}). This was followed by a decrease until the 1990s before stabilising within the most recent deposits (Cr = 39.0 ± 0.9 mg kg^{-1}; Mn = 140 ± 1 mg kg^{-1}; Ni = 17.0 ± 0.1 mg kg^{-1}; Cu = 67.2 ± 0.2 mg kg^{-1}; Zn = 279 ± 1 mg kg^{-1}; Cd = 3.34 ± 0.02 mg kg^{-1}) (Figure 6). Chromium, Mn, Ni, Cu, Zn and Cd share similar trends suggesting a common source for these TMM during the 1960s and 1970s. The maximum As concentration of 22.2 ± 0.2 mg kg^{-1} was measured at the bottom of the DAM17-02 core around 1945. Levels then gradually decreased and stabilised at 5.26 ± 0.14 mg kg^{-1} from the 1990s (Figure 6). Pb concentrations remained stable until the mid-1980s, then reached high concentrations during the 1990s–2000 with peaks of 854 ± 8 mg kg^{-1} in the early 1990s. Pb concentrations then declined from the mid-2000s and reached concentrations close to those measured at the beginning of the record. However, these were still higher than the lowest concentrations obtained in the late 1960s (Figure 6).

3.4. Comparison between geochemical background values, past and current concentrations

The total TMM concentrations measured by ICP-MS for the DAM17-02 core are also listed in Table 4. Vanadium and Th showed stable levels throughout the DAM17-02 core (V = 45.4 ± 5.4 mg kg^{-1} and Th = 5.18 ± 0.62 mg kg^{-1}), these were close to the GB values of the Eure River (V = 43.2 mg kg^{-1} and Th = 3.96 mg kg^{-1}; Table 2) and also close to the estimated levels for SPM (Figure 6). The temporal trends of V and Th confirm the absence of additional anthropogenic inputs within the watershed. Arsenic is the only TMM for which the lowest levels (corresponding to the most recent levels) were close to the GB value of 6.56 mg kg^{-1} (Table 2) and the levels measured on the SPM (Figure 6). The temporal trends in As confirm that there are no further anthropogenic additional inputs of As within the watershed. For the other TMM, the overall estimated concentrations in the DAM17-02 core were always higher than the GB values. Cobalt and Ag had similar trends to Cr, Mn, Ni, Cu, Zn or Cd and more precisely to Cr and Zn with a peak concentration in the 1970s (Figure 6). The most recent and oldest Cr, Co, Ni, Ag and Cd concentrations in the DAM17-02 core were close to those measured on the SPM. In contrast, for Cu and Zn, only the most recent concentrations were close to those measured on the SPM (Figure 6). Sb concentrations were similar to Pb with a peak concentration in the early 1990s of 6.62 mg kg^{-1}. The Pb levels measured in the SPM were slightly lower than those measured for the most recent deposits in the DAM17-02 core.

For all the TMM, the small differences in value between the estimated concentrations in SPM and those measured in most recent sedimentary deposits show that sediment cores collected were representative of SPM transfers in the Eure River.

3.5. Isotopic signature of Pb contamination

For the Seine Unit, where the Pb/Ti ratio was low (Figure S6), the values of the $^{206}Pb/^{207}Pb$, $^{206}Pb/^{204}Pb$ and $^{208}Pb/^{204}Pb$ ratios were higher than those recorded in the Eure Unit and lower

for $^{208}Pb/^{206}Pb$. The 1990s constituted a period of high Pb concentrations, where the $^{206}Pb/^{207}Pb$, $^{206}Pb/^{204}Pb$, $^{208}Pb/^{204}Pb$ ratios were low while the $^{208}Pb/^{206}Pb$ ratio was higher than the 1960s. This decrease in the $^{206}Pb/^{207}Pb$, $^{206}Pb/^{204}Pb$, $^{208}Pb/^{204}Pb$ ratios and increase in the $^{208}Pb/^{206}Pb$ ratio reflects a change in Pb source or inputs from a new source within the sediment record, contributing to the increased Pb levels. From the 2000s onwards, the $^{206}Pb/^{207}Pb$, $^{206}Pb/^{204}Pb$ and $^{208}Pb/^{204}Pb$ ratios increased slightly and $^{208}Pb/^{206}Pb$ remained stable. This increase was confirmed for the ratios calculated in 2012 ± 0.5 y while the $^{208}Pb/^{206}Pb$ started to decrease (Figure 5B). This increase and decrease for $^{208}Pb/^{206}Pb$, corresponds to the decrease in Pb concentrations recorded in the Martot Pond (Figure 5A). This reflects a reduction in the source contribution of Pb responsible for the elevated concentrations and the strong modification of the isotope ratios during the 1990s and 2000s. As such, there is a correlation between the temporal trend in Pb concentrations and that of Pb isotopes, the latter reflecting a major source change for Pb in the sediment record.

4. Discussion

4.1. Evaluation of anthropogenic contribution

The use of EFs to assess the anthropogenic impact of contamination requires the selection of a reference background. In some cases, the reference background used was a GB considered representative of the whole watershed, as for example for the Seine watershed (Horowitz et al., 1999; Thévenot et al., 2007; Le Cloarec et al., 2011), or the reference background was a GB measured at the bottom of sediment cores (Section 3.1.). In other cases, to compensate for the absence of a local GB, some studies have turned to the Upper Continental Crust (UCC) values (Wedepohl, 1995).To assess anthropogenic impacts, EFs were calculated in our study

474 from the DAM17-02 core using two reference backgrounds, the local GB defined using the
475 FOR17-02 core and a reference level corresponding to the lowest concentration at the bottom
476 or top of the DAM17-02 core (Table 5). Although the variations in EFs were similar
477 regardless of the reference background, the lowest EFs were those calculated using the
478 reference level (with the exception of As). The strongest EFs were in the 1960–1970s for Cr,
479 Co, Ni, Cu, Zn, Ag, Cd.

480 For Cr, Cu, Zn and Cd, the maximum EFs calculated from the GB of the Eure watershed,
481 were higher than those for the Oise River, a tributary of the Seine River (Le Cloarec et al.,
482 2011) or the Lot River (with the exception of Zn), a tributary of the Garonne River, known for
483 its historical Cd contamination with industrial origin (Audry et al., 2004). These EFs were
484 still lower than those calculated for the lower reaches of the Seine River, with the exception of
485 Cu and Cd (Le Cloarec et al., 2011) but were higher or close to those obtained for other major
486 French rivers such as the Rhône (Ferrand et al., 2012) or Loire rivers (Grosbois et al., 2012;
487 Dhivert et al., 2016). This demonstrates that with regard to these TMMs, anthropogenic
488 pressures within the Eure watershed were higher than those recorded for Oise River, another
489 tributary of the Seine estuary and of the same register as those recorded within major French
490 river watersheds. Conversely, the EFs calculated for As from the reference level were higher
491 but of the same order of magnitude, than those calculated from the GB on the watershed
492 (Table 5). Sediments from the Eure River had a maximum EF at the beginning of the
493 sedimentary record (i.e., during the 1940s), which was lower than the maximum EF calculated
494 in the Oise River or the lower reaches of the Seine River (Le Cloarec et al., 2011). In terms of
495 Sb, the maximum EF (Table 5) was lower than those recorded in the Oise River or the lower
496 reaches of the Seine River (Le Cloarec et al., 2011) whereas for Pb, the maximum EF in the
497 Eure River was higher than most maximum EFs obtained for other watersheds including the
498 Oise (Le Cloarec et al., 2011), Lot (Audry et al., 2004), Rhône (Ferrand et al., 2012) and

499 Loire rivers (Grosbois et al., 2012). This highlights the high anthropogenic pressures on the
500 Eure watershed since the 1990s, compared to other French watersheds. Since the 1990s, all of
501 the TMM (except for Sb and Pb) were stable with relatively low concentrations (and EFs),
502 which may result from a significant reduction or a halt in additional discharges during the
503 1960–1970s period. This is assuming that a strong EF corresponds to strong anthropogenic
504 pressure. The lowest levels, considered the baseline within the sediment record, remain higher
505 compared to the GB of the watershed, with the exception of As. This suggests that
506 anthropogenic pressures within the watershed are now low but not null (Table 5). Due to the
507 multiple anthropogenic pressures that a watershed may experience, it can be assumed that soil
508 and sediments quality can only be degraded to pre-industrial revolution levels. As such, EF
509 from the bottom or top of sediment cores, which can be considered permanent anthropogenic
510 baseline, may seem more realistic for the evaluation of anthropogenic pressures in a
511 watershed. This makes it possible to distinguish between periods when anthropogenic
512 pressure is active via industrial, agricultural or even urban discharges, and periods when high
513 levels of TMM (in relation to GB) are solely due to the mechanical erosion of impacted soils
514 or re-suspension of contaminated sediments stored in the river. The remobilisation of
515 contaminated soils has already been observed in small wine-growing watersheds (Sabatier et
516 al., 2014) or those dominated by pastoralism (Bajard et al., 2018).
517
518 4.2. Current and theoretical past contributions
519
520 Over the period 1994–2003, Thévenot et al. (2007) estimated mean annual flux from the
521 Seine watershed to the Seine estuary for Cr, Cu, Zn, Cd and Pb. Assuming that TMM flux out
522 of the Seine watershed has not changed since this period, this flux out of the Eure watershed
523 for 2017 (Table 3) would be representative of 7, 8, 9, 10, and 16% of the total inputs (Seine

524 River + Eure River) of Cr, Cu, Zn, Cd and Pb into the Seine estuary, respectively (Figure 7).
525 It is also interesting to estimate the magnitude of inputs from the Eure watershed in the 1990s,
526 a period when very high Pb concentrations were recorded. Streamflow in the Eure River has
527 been near constant since 1971 (www.hydro.eaufrance.fr), assuming that 2017 corresponds to a
528 medium hydrological cycle in terms of hydro-sedimentary inputs. As such, theoretical flux
529 estimation for previous years using the SPM flux for the year 2017 and annual TMM
530 concentrations within sediment cores could prove insightful. The average Pb concentration
531 over 1992–1996 was 846 mg kg^{-1} (Figure 6; Table 3), which equates to an average flux of 95 t
532 y^{-1}, making the Eure watershed the main contributor of Pb to the Seine estuary (52%) (Figure
533 7). This confirms the river's classification as very high contamination by Meybeck et al.
534 (2004), and that the watershed has recently experienced major Pb contamination with little
535 national equivalency.
536
537 4.3. Sources of contamination
538
539 Similar temporal trends for Cr, Co, Ni, Cu, Zn, Ag, Cd (Figures 5A and 6) suggest that a
540 single anthropogenic source was responsible for the high levels over the 1960s–1970s. The
541 presence of Cd, Ni, Zn could be related to the activities and discharges of a saline batteries
542 factory. Within the Eure watershed there was a saline batteries factory (opened in the late
543 1930s), supplying most of the national market during the 1960s, resulting in very high
544 localised production. The factory was the subject of an increased administrative survey by a
545 national agency in the early 1970s (Agence Financière de Bassin “Seine-Normandie,” 1971).
546 This survey, coupled with the development of alkaline batteries (manufactured abroad), may
547 have caused a drop in production, resulting in a reduction in concentrations after the 1970s,
548 until the closure of the industrial site in 1994 (Gardes et al., 2020a).

549 The main source responsible for As release has not been defined but may be attributed to
550 different anthropogenic activities. This includes discharge from agricultural activities (e.g.,
551 wood preservative), industrial activities (glass and electronics industry, metal treatment,
552 ammunition manufacturing, production of dyes and colorants or metal smelters smelting) or
553 coal combustion (Reimann et al., 2009). But it must be noted that arsenical herbicides have
554 been forbidden in France since 1973; it is probable that it was one of the main sources in the
555 Eure watershed as the current concentrations, near the GB value, were reached after the
556 1970s.
557 Sources responsible for Pb contamination can be discriminated from Pb isotopes,
558 particularly in tri-isotopic diagrams in which these potential sources act as end-members (e.g.,
559 Cloquet et al., 2015). In the diagram $^{208}Pb/^{206}Pb$ versus $^{206}Pb/^{207}Pb$, the measurement within
560 the Seine Unit U2 ($^{208}Pb/^{206}Pb$ = 2.070 and $^{206}Pb/^{207}Pb$ = 1.193) was close to the points with a
561 natural (or pre-industrial) signature ($^{208}Pb/^{206}Pb$ = 2.053 and $^{206}Pb/^{207}Pb$ = 1.2007, Elbaz-
562 Poulichet et al., 1986; $^{208}Pb/^{206}Pb$ = 2.066 and $^{206}Pb/^{207}Pb$ = 1.197, Monna et al., 1997;
563 $^{208}Pb/^{206}Pb$ = 2.079 and $^{206}Pb/^{207}Pb$ = 1.200, Ferrand et al., 1999) (Figure 8A). Moreover, such
564 a high $^{206}Pb/^{207}Pb$ ratio has never been previously recorded in the Seine watershed (Ayrault et
565 al., 2012) with the exception of the upstream Yonne River where $^{206}Pb/^{207}Pb$ was
566 approximately 1.19 prior to the 1960s and after the 2000s. This reflects an absence of
567 anthropogenic Pb contamination (Ayrault et al., 2010). In terms of the Eure Unit U1, the
568 signature measured in 1965 ($^{208}Pb/^{206}Pb$ = 2.097 et $^{206}Pb/^{207}Pb$ = 1.170) showed a mix
569 between pre-industrial and industrial poles. It is possible to assess the relative contribution of
570 a source using the binary model, described for example by Komárek et al. (2008). Based on
571 the $^{206}Pb/^{207}Pb$ ratio proposed by Monna et al. (1997) as the average signature of French
572 industrial Pb ($^{206}Pb/^{207}Pb$ = 1.155), the contribution of the unidentified industrial source, was
573 in the range 60–67%. This is dependent on the value used for the pre-industrial source,

574 between the Seine Unit from this study, and the values proposed by Elbaz-Poulichet et al.
575 (1986), Monna et al. (1997) and Ferrand et al. (1999). This shows that the contribution of
576 anthropogenic Pb was already significant during the 1960s in the Eure watershed, whereas Pb
577 concentrations recorded in sediment cores at the same time were not high and did not exhibit
578 an increasing trend (Figure 5A). However, the estimation of this contribution does not
579 consider a possible existing contribution from other sources such as coal use (e.g., coal power
580 plant, coal combustion), which is also a source of Pb. Notably, the isotope signature recorded
581 in 1965 was close to the isotope signatures of Pb from coal use (Cloquet et al., 2006; Elbaz-
582 Poulichet et al., 2011). Although it is unlikely that there existed only a single source emitting
583 Pb during the 1960s. This 1960s signature was close to the European Standard Lead Pollution
584 (ESLP) line defined by Haack et al. (2002, 2003) from atmospheric Pb deposition across
585 Europe, confirming the presence of anthropogenic Pb in this sample. More recent
586 contamination signatures (1995, 1997, 2002 and 2012) were located between the industrial
587 and leaded petrol poles (defined from data from Monna et al. (1997), Véron et al. (1999) and
588 Cloquet et al. (2015)). These were aligned with the industrial Pb line also defined by Haack et
589 al. (2002, 2003) from industrial equipment, exhaust pipes and several major ore deposits. As
590 such, it is possible to envisage a mixture between these two poles to understand the signatures
591 obtained recently in the Eure watershed. Figure 8B shows that these signatures, in particular
592 that of 1995, were close to a Pb smelting signature defined in northern France from aerosols
593 (n = 2; $^{208}Pb/^{206}Pb = 2.1269 \pm 0.0008$ et $^{206}Pb/^{207}Pb = 1.1330 \pm 0.0009$ (Véron et al., 1999)),
594 or dust from a Pb-Zn smelting plant (n = 2; $^{208}Pb/^{206}Pb = 2.124 \pm 0.002$ et $^{206}Pb/^{207}Pb =$
595 1.1289 ± 0.0023 (Cloquet et al., 2006)) (not shown in Figure 8B). However, these authors do
596 not consider contributions from leaded gasoline. Therefore, the signature recorded here for
597 these four points was the result of an anthropogenic Pb input, the partial source of which
598 could be the same as that recorded by the signature measured in 1965. The difference in

599 signature between 1965 and more recent signatures may also be attributable to an increase in a
600 major anthropogenic source not yet identified but unrelated to coal (indicated by increases in
601 Pb concentrations as seen in the sediment cores). There is an anthropogenic source of Pb in
602 the Eure watershed related to industrial activities that commenced in the 1960s and increased
603 from the 1990s. Chiffoleau et al. (2012) used a sediment core sampled in the Seine estuary to
604 demonstrate that there was a single anthropogenic source which accounts for the variations in
605 the $^{206}Pb/^{207}Pb$ ratio in the estuary since 1975. Between 1975 and 1985 the variations in this
606 ratio are explained by a mixture between intra-estuarine phosphogypsum discharges and an
607 unknown anthropogenic source Post-1985 the variations may be attributed to a mixture
608 between a natural source and this same unknown anthropogenic source. The isotopic signature
609 of this source ($^{208}Pb/^{206}Pb$ = 2.132 et $^{206}Pb/^{207}Pb$ = 1.13) was isolated by Chiffoleau et al.
610 (2012) and Figure 8B shows that it is close to the anthropogenic signature recorded by the
611 Eure River sediments, indicating that it is likely the source comes from the Eure River. Based
612 on the SPM sampled in the Eure River in March and September 2009 (and of dissolved Pb
613 (Pb_d) measurements), Chiffoleau et al. (2012) estimated a low $^{206}Pb/^{207}Pb$ ratio, in the range
614 1.12–1.14, consistent with ratios from our study (Figure 8B). From a sampling campaign
615 conducted in September 2011, Chiffoleau et al. (2012) found an increase in Pb_d and a
616 $^{206}Pb/^{207}Pb$ ratio of about 1.13 downstream of the city of Dreux (and a $^{206}Pb/^{207}Pb$ ratio greater
617 than 1.14 upstream). Assuming the mixture of natural and anthropogenic sources, the authors
618 estimated that the additional anthropogenic source had a $^{206}Pb/^{207}Pb$ signature of
619 approximately 1.12. A cathode-ray tube (CRT) factory began operation in Dreux in 1956 and
620 a second factory, specialising in the assembly of CRT TVs and the manufacture of electrical
621 circuits, opened in 1974 at the same industrial site. Chiffoleau et al. (2012) hypothesised that
622 the CRT factory represented this unknown Pb source. CRT may contain up to 3 kg of Pb
623 (Tsydenova and Bengtsson, 2011). This factory, as well as the CRT TVs assembly factory,

624 was purchased in the early 1990s to boost the economy (and thus increase CRT production)
625 prior to closing in 2006 and 2010 for the CRT factory and CRT TVs assembly factory,
626 respectively. In 2001, Pb discharge to water from the industrial site was ranked among the
627 highest in Europe by the European Environment Agency (E-PRTR, 2006). Increases in Pb_p
628 within sediment cores since the 1990s, as well as isotope signatures measured in 1995, 1997
629 and 2002, reflect the economic recovery initiated by this industry. However, the high
630 concentrations measured from the end of the 2000s in the sediment cores, and the isotopic
631 signature estimated in 2012, demonstrate that anthropogenic Pb inputs are still present in the
632 Eure watershed despite the cessation of industrial activities. These inputs are sufficiently high
633 to modify the $^{206}Pb/^{207}Pb$ ratio in the SPM of the Seine estuary as shown by Chiffoleau et al.
634 (2012).
635
636 **5. Conclusions**
637
638 This study of sediment cores has revealed that the lower reaches of the Eure watershed have
639 been subject to significant anthropogenic pressures since the 1940s, at comparable or higher
640 levels, than those experienced by other French watersheds during the twentieth century. The
641 temporal trends of TMM concentrations have shown strong anthropogenic impacts, mainly of
642 industrial origin during the 1960s and 1970s. A general trend towards a decrease in impacts
643 was emerging at the end of the twentieth century, as in the majority of French river
644 watersheds. However, an additional source, potentially of industrial origin, was responsible
645 for high Pb discharges that heavily contaminated sediments of the Eure River. It had a
646 signature found in sediments of the Seine estuary, that had already been heavily impacted by
647 inputs from the fluvial part of Seine River, one of the most anthropized catchments in Europe.
648 The similarity between the levels of TMM in the most recent sedimentary deposits and on the

SPM showed that the sedimentary records were representative of SPM transfers, and that the footprint of industrial contamination was still present despite the cessation of the activities responsible for the discharges. This footprint reflects the existence of permanent anthropogenic baseline within the river, resulting from the erosion of contaminated soils and sediments stored in the river and remobilised during floods and/or landscaping projects. As a result, the current flux in TMM to the Seine estuary remain significant despite the cessation or drastic reduction of discharges from anthropogenic activities. Based on the estimation of past theoretical flows, it has been shown in the case of Pb that the Eure watershed could exceed contributions from the Seine watershed, making the former the main contributor of Pb to the Seine estuary during the 1990s.

Reference data

The dataset is available at https://doi.org/10.1594/PANGAEA.912843 (Gardes et al., 2020b) and at https://doi.org/10.1594/PANGAEA.911102 (Gardes et al., 2020c).

Acknowledgements

This work is part of the OSS 276 project, financially supported by the Seine-Normandie Water Agency (France). This work is support by the CNRS toward EC2CO grant (Avant-Seine project). Gardes Th. grant was funded by the Region Normandie, which supports the scientific consortium SCALE UMR CNRS 3730. CNRS also supports the French national cyber-Repository (https://www.cybercarotheque.fr/) which is a portal for metadata associated with marine, lake/river sediment cores stored in French laboratories. This project was initiated by CLIMCOR project (http://climcor-equipex.dt.insu.cnrs.fr/).

Our paper is a contribution to ROZA-LTER-fr project : Retro-Observatory of sedimentary Archive from French LTER (http://ccwbvps18.in2p3.fr/maps/visualiseur-coyoxhup#project) which is a portal for data associated with marine, lake sediment cores stored in LTER French laboratories.

Authors sincerely thank the “Parc Naturel Régional du Perche” for allowing the access at the ponds and Aurélie Tran Van Loc for her participation at the coring session.

Authors also thank Romain Levaillant and Michel Simon (M2C laboratory) for their participation in the field sampling (coring sessions, TIMS deployment).

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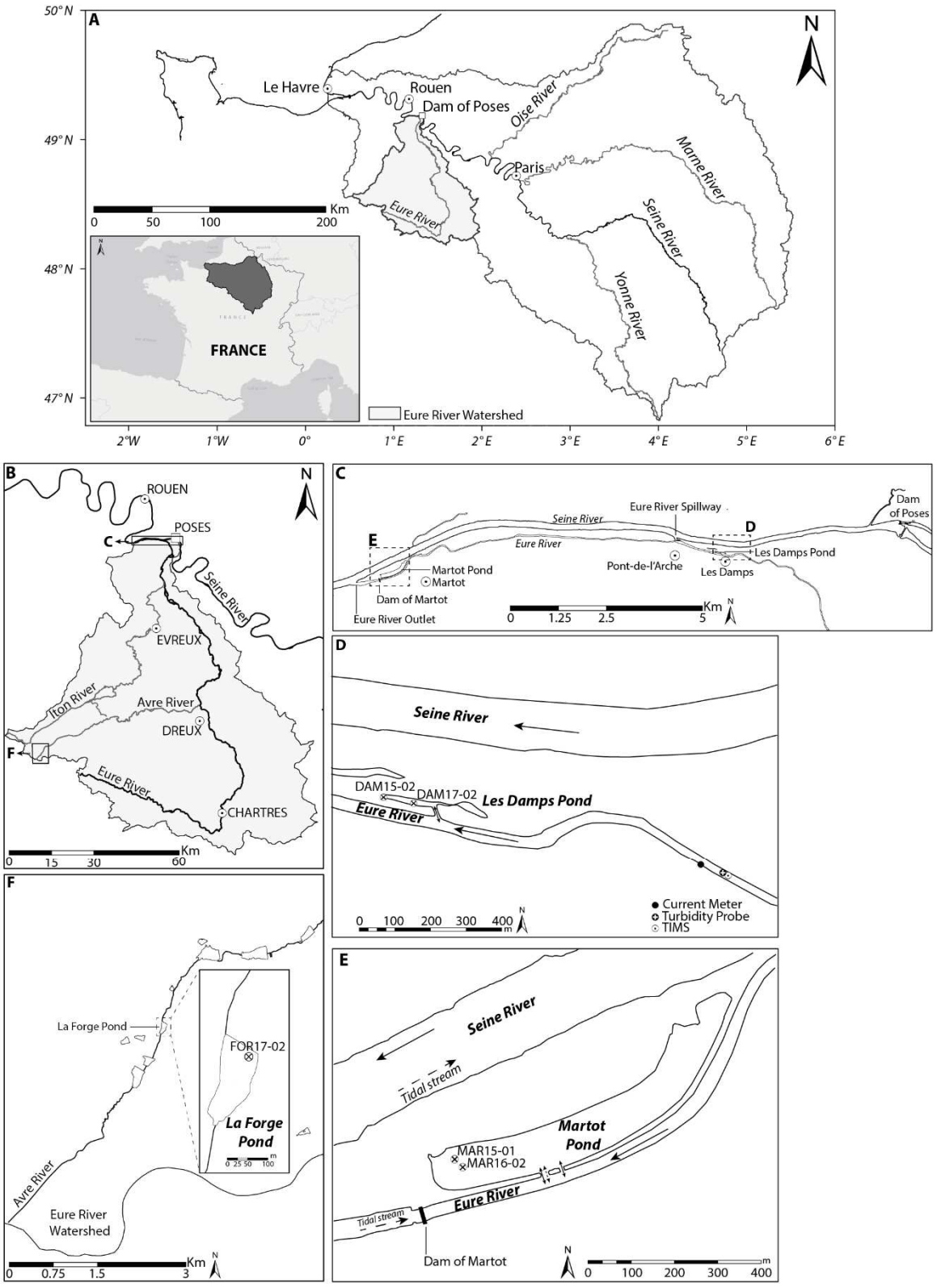


Figure 1. A. The Seine River Watershed; B. The Eure River Watershed; C. Study Area; D. Sediment core locations in the Les Damps Pond and current meter, turbidity probe and TIMS locations; E. Sediment core locations in the Martot Pond and F. Sediment core location in the La Forge Pond.

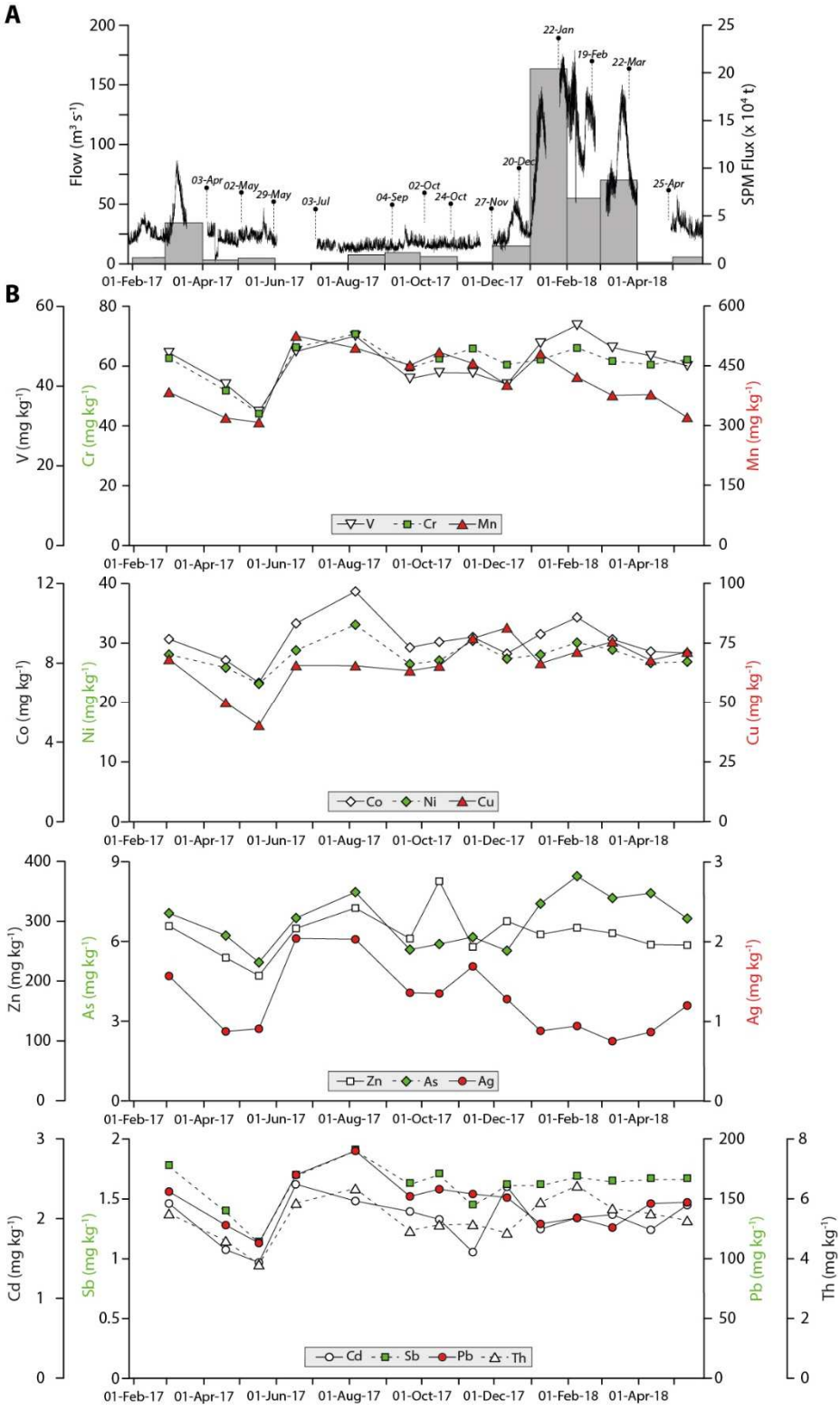


Figure 2. A. Flow (m^3 s^{-1}) and SPM Flux (x 10^4 t $month^{-1}$) at the Les Damps site; B. TMM concentrations (mg kg^{-1}) at the Les Damps site.

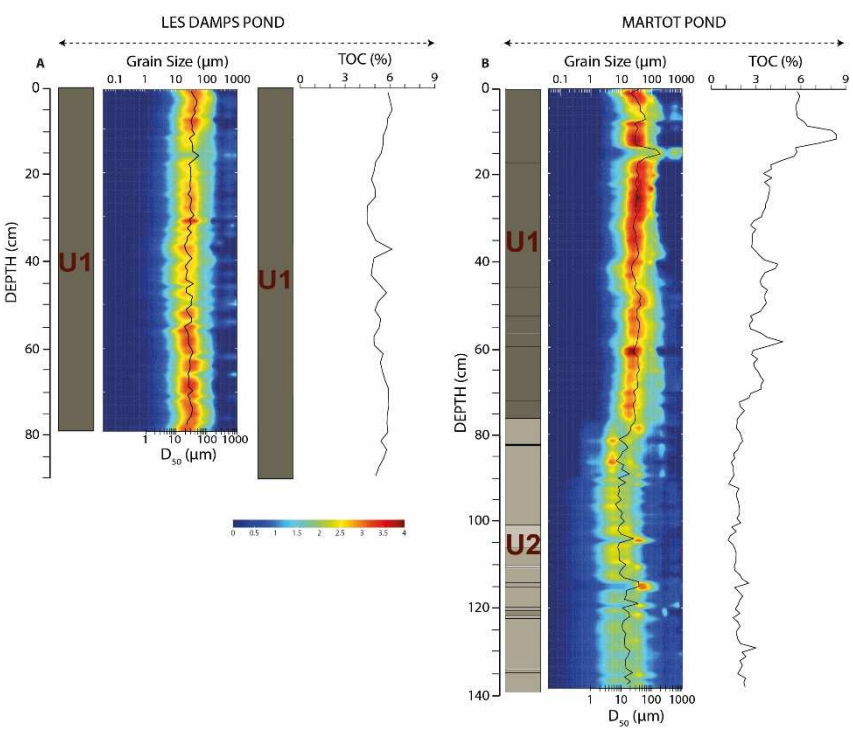


Figure 3. A. Grain size distribution, D_{50} along the DAM15-02 core and TOC along the DAM17-02 core; B. Grain size distribution, D_{50} and TOC along the MAR15-01 core.

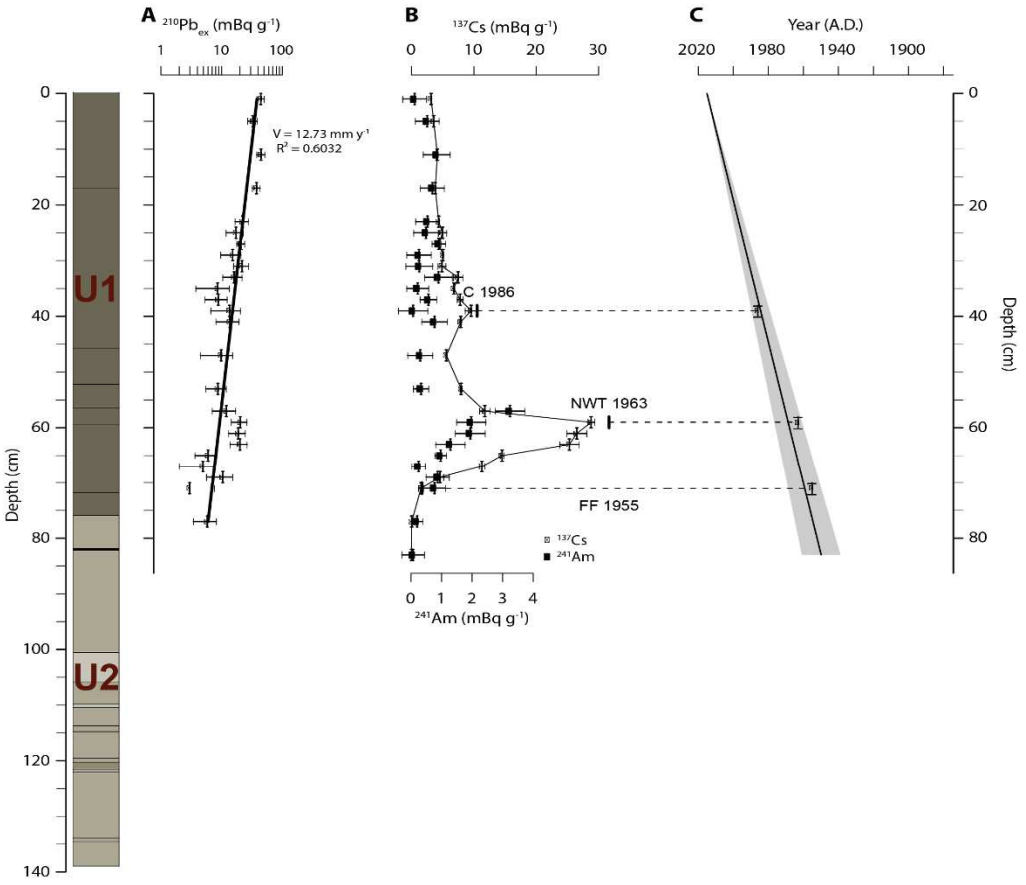


Figure 4. Age model for MAR15-01 core building with A. $^{210}Pb_{ex}$; B. ^{137}Cs and. ^{241}Am (the first identification of ^{137}Cs in 1955 (FF 1955), the maximum fallout from atmospheric nuclear weapon testing in 1963 (NWT 1963) and the Chernobyl accident in 1986 (C 1986) are represented); and C. Age model.

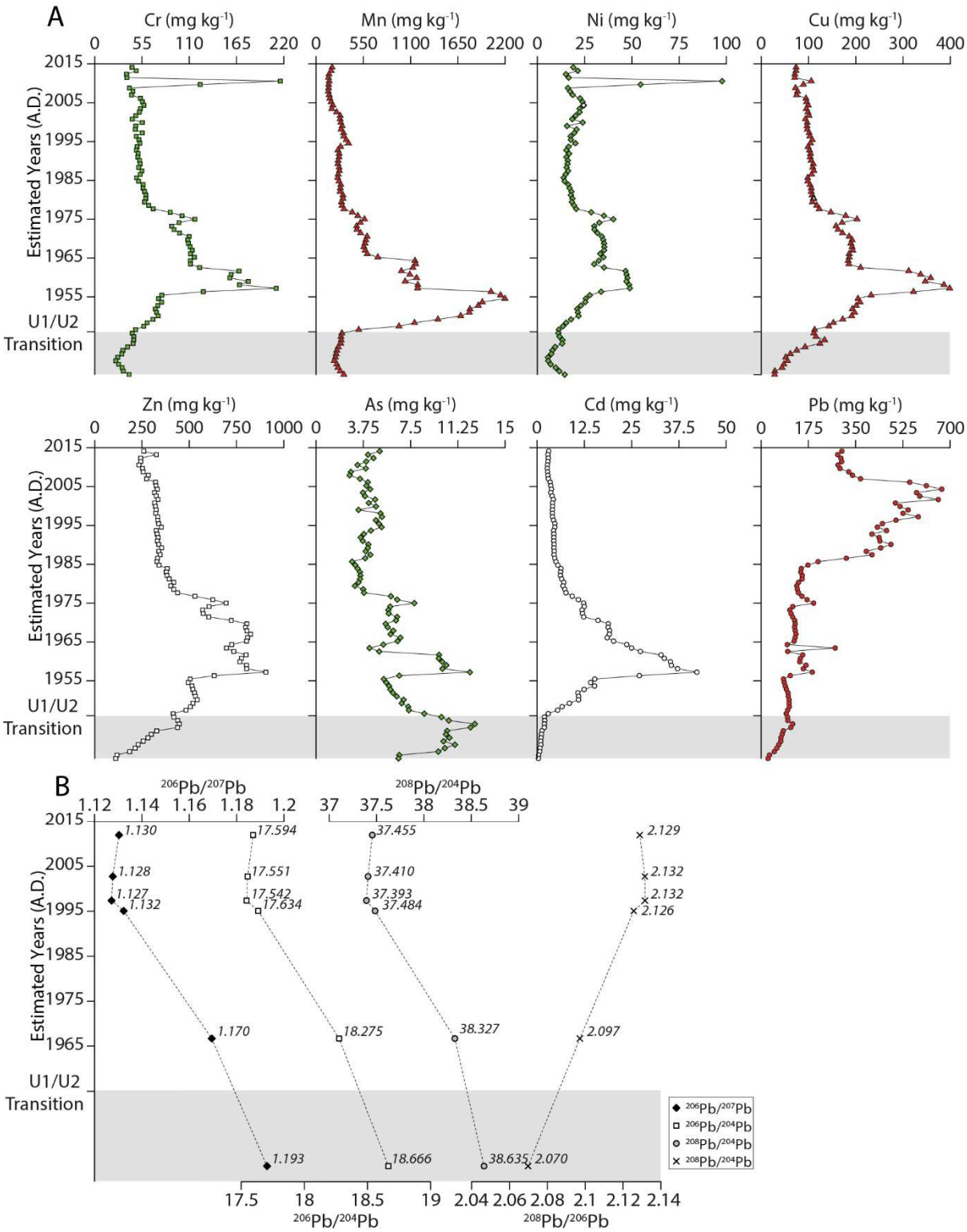


Figure 5. A. TMM concentrations along the MAR16-02 core; B. Pb isotope ratios along the MAR15-01 core (the Seine Unit U2 is represented in grey).

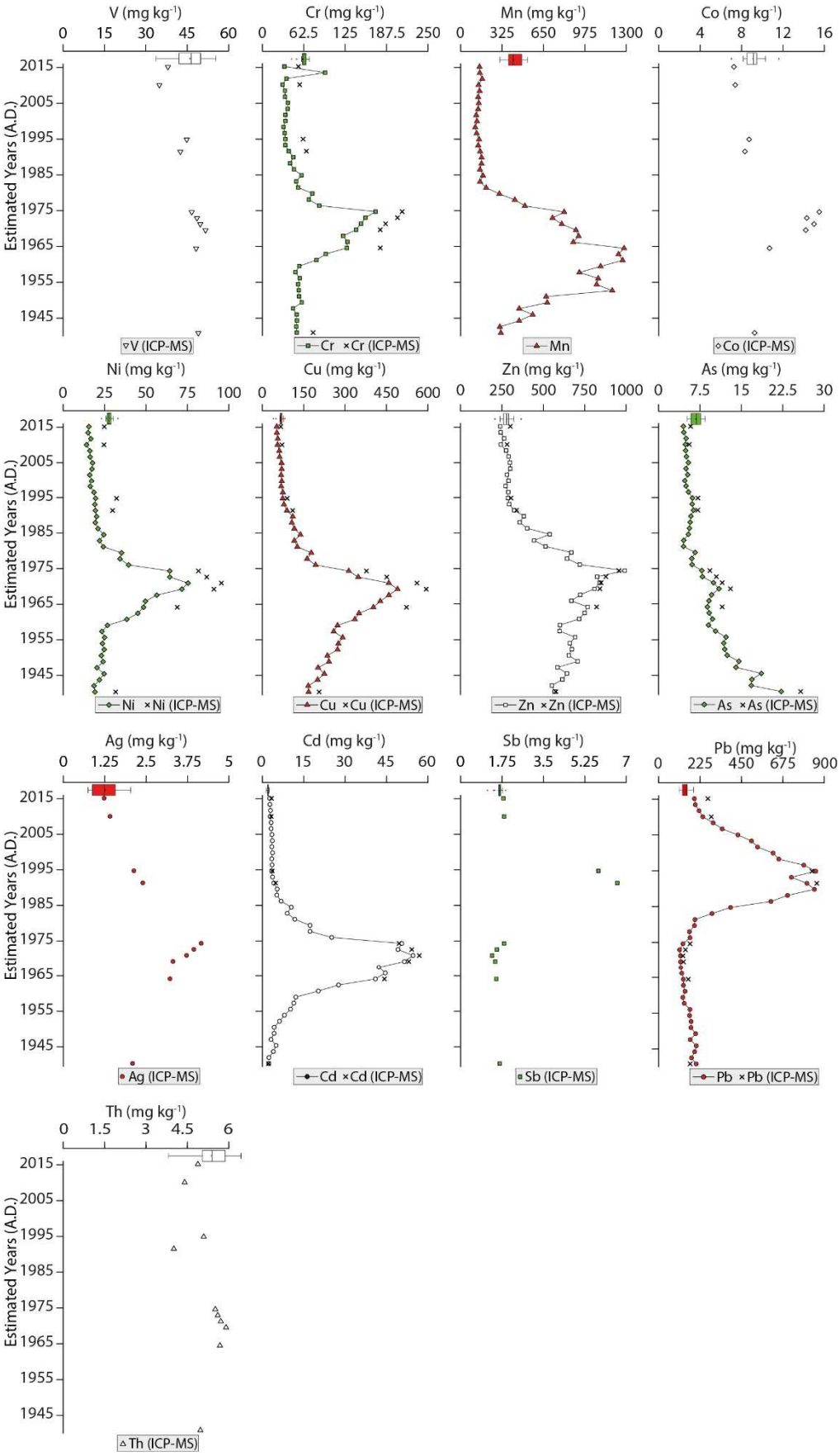


Figure 6. TMM concentrations along the DAM17-02 core. Boxplot representing TMM concentrations from SPM.

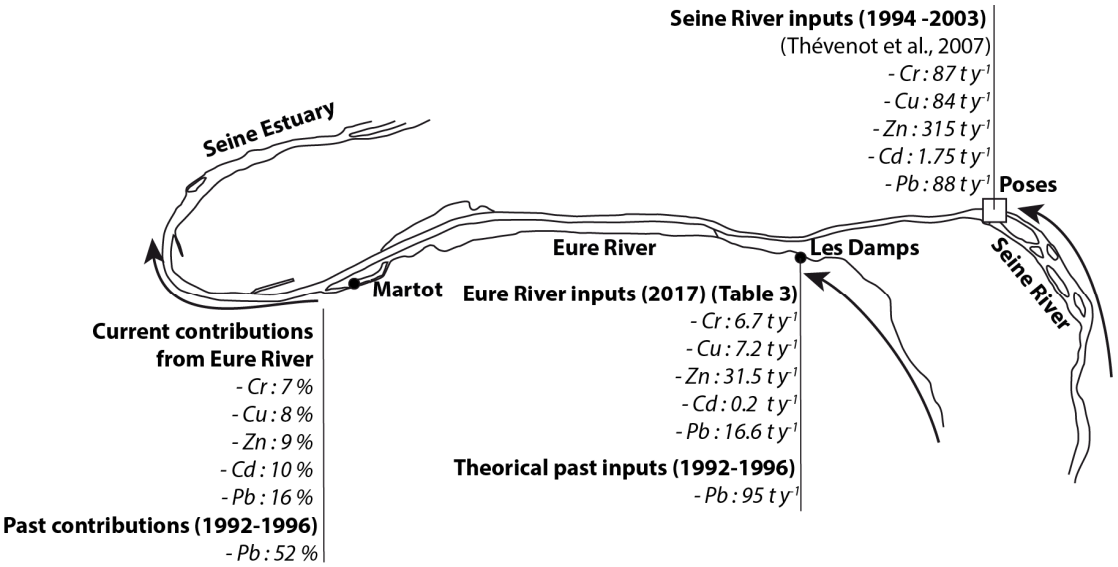


Figure 7. Eure River inputs to the Seine estuary and estimation of Eure River contributions.

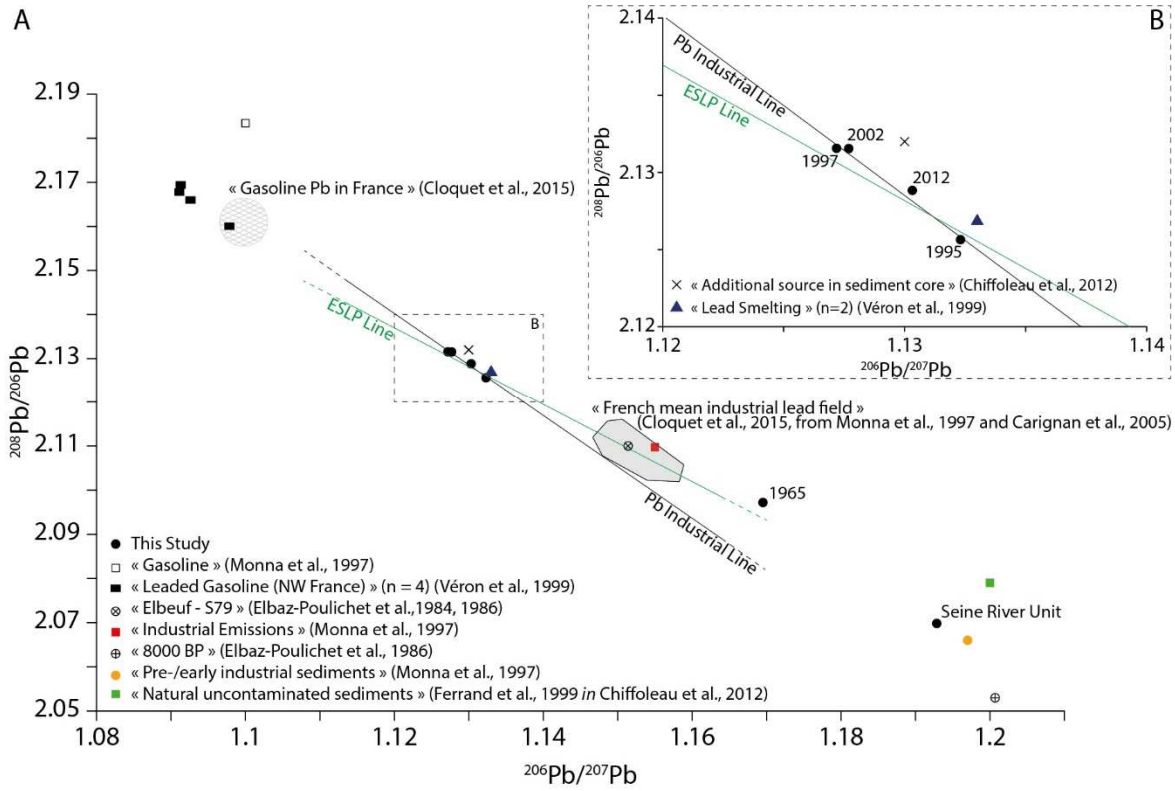


Figure 8. Three isotope diagrams $^{208}Pb/^{206}Pb$ versus $^{206}Pb/^{207}Pb$.

Table 1. Characteristics of sediment cores collected in the La Forge Pond, the Les Damps Pond and the Martot Pond (WGS 84).

Pond	Core ID	IGSN	Longitude (X)	Latitude (Y)	Core Length (cm)
LES DAMPS	DAM15-02	IEM2C0016	1°10'9.05" E	49°18'16.13" N	80
	DAM17-02	IEM2C000E	1°10'13.26" E	49°18'15.66" N	90
MARTOT	MAR15-01	IEM2C0001	1°03'1.68" E	49°17'49.68" N	138
	MAR16-02	IEM2C0008	1°03'2.60" E	49°17'49.30" N	129
LA FORGE	FOR17-02	IEM2C001D	0°37’47.89" E	48°37’51.75" N	69.5

Table 2. Age (A.D.), D_{50} (µm), TOC (%) and TMM (mg kg^{-1}) at the bottom of the FOR17-02 core.

AGE	D_{50}	TOC	V	Cr	Co	Ni	Cu	Zn	As	Ag	Cd	Sb	Pb	Th
1646 - 1669	20.9	2.77	43.2	32.5	5.58	9.78	5.60	32.3	6.56	0.19	0.11	0.41	10.6	3.96

Table 3. Flux of particulate TMM (kg month^{-1}) at the Les Damps site.

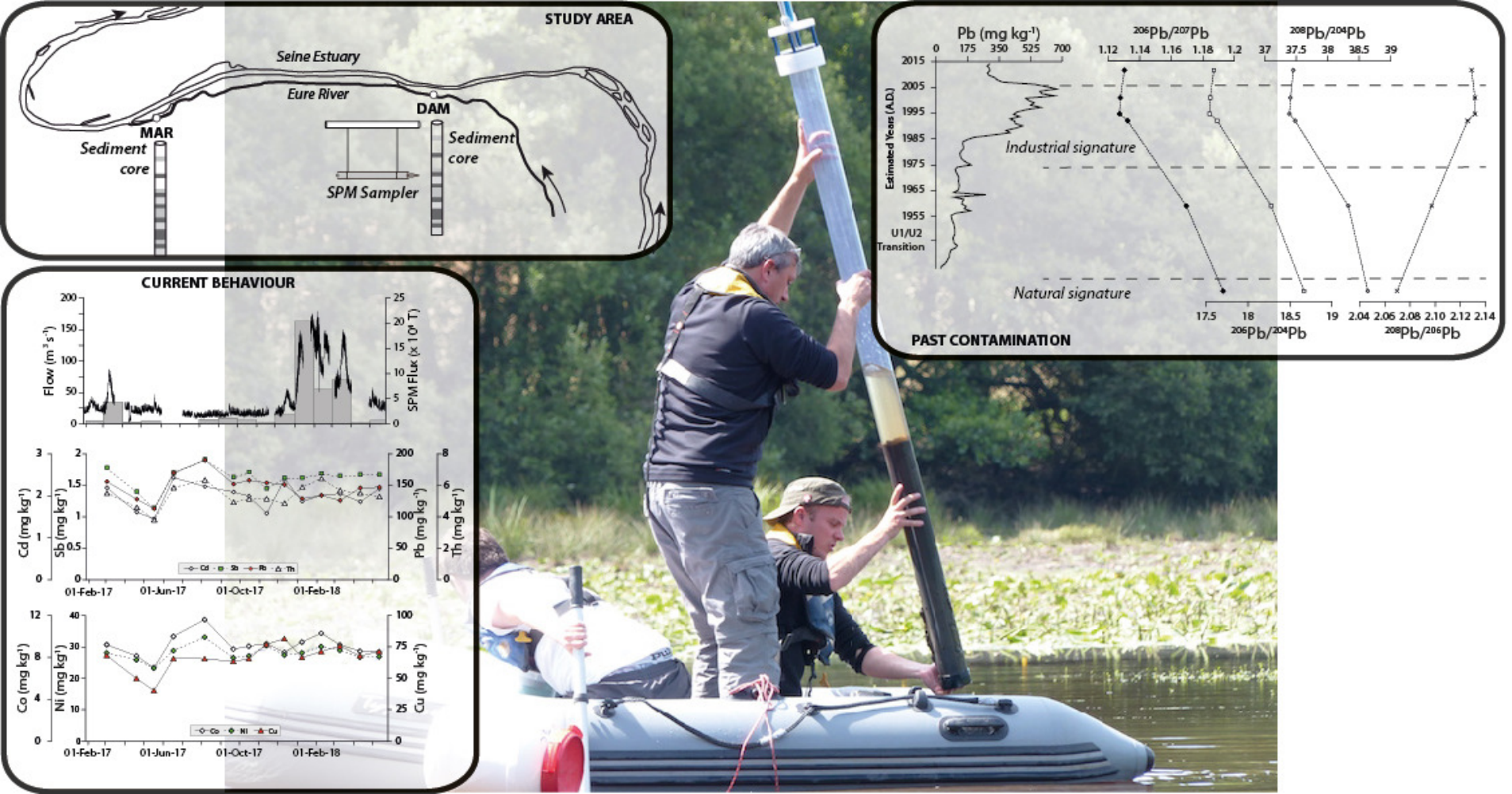
Sampling Interval	V	Cr	Mn	Co	Ni	Cu	Zn	As	Ag	Cd	Sb	Pb	Th
27-Jan-17 – 31-Jan-17	11.3	14.6	89.2	2.14	6.55	15.9	68.0	1.64	0.36	0.51	0.41	36.3	1.28
Feb-17	300	389	2378	57.1	175	423	1812	43.8	9.73	13.6	11.0	967	34.1
Mar-17	2065	2676	16359	393	1201	2908	12462	301	66.9	93.3	75.5	6651	235
Apr-17	166	212	1307	33.4	106	204	980	25.5	3.58	6.60	5.71	523	18.8
May-17	199	261	1830	41.3	135	240	1226	30.6	5.54	8.65	6.68	669	22.6
Jun-17	-	-	-	-	-	-	-	-	-	-	-	-	-
Jul-17	42.7	57.5	402	9.40	26.9	53.2	261	6.37	1.65	1.81	1.54	155	5.16
Aug-17	484	651	4549	106	304	602	2957	72.2	18.6	20.4	17.5	1751	58.4
Sep-17	497	704	5344	104	314	750	3214	67.5	16.2	24.8	19.2	1800	59.8
Oct-17	328	475	3634	68.8	208	504	2654	44.9	10.5	14.8	12.6	1190	39.0
Nov-17	74.5	113	782	16.2	52.3	134	450	10.6	2.87	2.85	2.52	266	8.88
Dec-17	843	1147	8165	167	519	1402	5437	120	20.7	40.6	30.1	2646	99.5
$\sum_{2017}$	5010	6701	44839	999	3047	7236	31521	724	157	228	183	16653	582
Jan-18	10904	13155	90856	2028	5972	14137	58065	1637	187	398	336	26908	1267
Feb-18	3757	4505	28591	698	2058	4926	19793	574	63.4	138	115	9138	436
Mar-18	4332	5390	32852	801	2511	6568	24403	669	66.6	178	144	11131	495
Apr-18	*73.4*	*101*	*522*	*13.9*	*43.8*	*116*	*424*	*11.2*	*1.95*	*3.54*	*2.70*	*240*	*8.60*
01-May-18 – 22-May-18	304	419	2160	57.3	181	481	1755	46.2	8.08	14.7	11.2	992	35.6

Table 4. TMM concentrations (mg kg^{-1}) in the DAM17-02 core (ICP-MS).

Sampling Depth (cm)	V	Cr	Co	Ni	Cu	Zn	As	Ag	Cd	Sb	Pb	Th
1	38.0	54.4	7.26	24.8	67.1	300	5.81	1.23	3.34	1.81	268	4.88
7	34.9	56.3	7.39	24.6	71.6	281	5.58	1.41	3.24	1.84	286	4.40
25	44.8	61.5	8.73	32.2	90.2	305	7.13	2.13	3.61	5.82	833	5.09
29	42.5	66.5	8.32	29.8	110	339	7.13	2.40	4.84	6.62	859	4.01
49	46.6	211	15.5	81.6	377	956	9.31	4.16	49.6	1.84	171	5.51
51	48.5	204	14.3	86.6	451	877	10.47	3.94	54.1	1.53	146	5.60
53	49.7	186	15.0	95.5	560	849	11.54	3.72	56.9	1.33	133	5.71
55	51.6	178	14.2	91.0	594	841	13.03	3.31	53.1	1.46	135	5.90
61	48.2	178	10.7	68.9	523	821	11.54	3.22	44.2	1.51	161	5.68
89	49.0	76.6	9.26	31.7	206	576	25.74	2.09	2.25	1.65	171	4.97

Table 5. Enrichment Factor (EF) based on the reference level from the DAM17-02 core (EF_{DAM}) and the geochemical background of the Eure River Watershed (EF_{FOR}).

	Cr		Co		Ni		Cu		Zn		As		Ag		Cd		Sb		Pb	
Depth (cm)	EF_{DAM}	EF_{FOR}	EF_{DAM}	EF_{FOR}	EF_{DAM}	EF_{FOR}	EF_{DAM}	EF_{FOR}	EF_{DAM}	EF_{FOR}	EF_{DAM}	EF_{FOR}	EF_{DAM}	EF_{FOR}	EF_{DAM}	EF_{FOR}	EF_{DAM}	EF_{FOR}	EF_{DAM}	EF_{FOR}
1	1.0	1.4	1.00	1.1	1.0	2.1	1.0	9.7	1.0	7.5	1.0	0.7	1.0	5.2	1.5	24	1.1	3.6	1.6	21
7	1.1	1.6	1.1	1.2	1.1	2.3	1.2	12	1.0	7.8	1.1	0.8	1.3	6.6	1.6	26	1.3	4.0	1.9	24
25	1.1	1.5	1.2	1.2	1.2	2.6	1.3	13	1.0	7.3	1.2	0.8	1.7	8.7	1.6	25	3.4	11	4.7	61
29	1.5	2.0	1.4	1.5	1.5	3.0	2.0	20	1.4	10	1.5	1.1	2.4	12	2.7	42	5.0	16	6.2	80
49	3.4	4.7	1.9	2.0	2.9	6.0	5.0	48	2.8	21	1.4	1.0	3.0	16	20	313	1.0	3.2	0.9	12
51	3.3	4.4	1.7	1.8	3.0	6.3	5.9	57	2.5	19	1.6	1.1	2.8	15	21	336	0.8	2.6	0.8	10
53	2.9	4.0	1.8	1.9	3.3	6.8	7.1	69	2.4	18	1.7	1.2	2.6	14	22	346	0.7	2.2	0.7	8.7
55	2.7	3.7	1.6	1.7	3.0	6.2	7.3	71	2.3	17	1.9	1.3	2.2	12	20	312	0.7	2.4	0.7	8.6
61	2.8	3.8	1.3	1.3	2.4	4.9	6.7	65	2.3	18	1.7	1.2	2.2	12	17	270	0.8	2.6	0.8	11
89	1.4	1.9	1.3	1.3	1.3	2.6	3.0	29	1.9	14	4.4	3.1	1.7	9	1.0	16	1.0	3.2	1.0	13


STUDY AREA
Seine Estuary
Eure River
DAM
MAR
Sediment core
Sediment core
SPM Sampler
CURRENT BEHAVIOUR
Flow (m³ s⁻¹)
SPM Flux (x 10⁴ T)
Cd (mg kg⁻¹)
Sb (mg kg⁻¹)
Pb (mg kg⁻¹)
Th (mg kg⁻¹)
Cd Sb Pb Th
01-Feb-17
01-Jun-17
01-Oct-17
01-Feb-18
Co (mg kg⁻¹)
Ni (mg kg⁻¹)
Cu (mg kg⁻¹)
Co Ni Cu
01-Feb-17
01-Jun-17
01-Oct-17
01-Feb-18
Pb (mg kg⁻¹)
206Pb/207Pb
208Pb/204Pb
Estimated Years (A.D.)
U1/U2 Transition
Industrial signature
Natural signature
206Pb/204Pb
208Pb/206Pb
PAST CONTAMINATION